

*APPLICATION OF PHOTOMETRIC ANALYSES FOR INTERPRETATION  
OF MARS-MARINER VI AND VII IMAGERY*

Author: Robert E. Kinzly

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## SECTION 1

### INTRODUCTION AND SUMMARY

This is the final report under Contract NASW-2222 sponsored by the Planetary Geology Program at NASA Headquarter (Code SL). The objective of this program was to investigate and illustrate the application of radiometric analyses in the interpretation of Mariner VI and VII imagery. We felt that the following types of problems could be addressed using this imagery:

- (1) Evaluation of local reflectivity changes,
- (2) Augmentation of geologic mapping,
- (3) Discrimination of atmospheric phenomena, and
- (4) Investigation of polar cap structure.

The use of conventional radiometric techniques requires the ability to convert sensor response into surface radiance, that is radiometric calibration of the Mariner TV systems. In order to interpret the surface radiance, knowledge of the bi-directional reflectivity function and the atmospheric effects (if any) are required. To circumvent the requirement that these functions, i.e., the radiometric response and bi-directional reflectivity, be known with a high degree of accuracy an interpretation technique based upon reflectivity ratios was employed during this study. Alternate near encounter frames from the wide angle Mariner cameras were taken through different spectral filters (red, green or blue). Consequently, the overlap between successive frames show the same area of the martian surface in two spectral bands. We call these areas the bi-band coverage areas.

A comprehensive review of previous and subsequent efforts to the present study was made and is included in this report. In performing this survey it was obvious to us that the terminology and definitions employed in discussing radiometric investigations (both inter-planetary and earth-based) is quite divergent and confusing to the reader. This lack of precision makes the comparison of results difficult. Consequently the next section of this report proposes a radiometric model which incorporates the definitions and terms used in both astronomy and electro-optics.

As noted above, the interpretation of radiometric data can require knowledge of the bi-directional reflectivity or photometric function. Consequently in Section 2 we have reviewed the definition of this function and discuss several parametric models for its representation. We also have included the parametric values obtained by several authors in fitting observed radiometric data to the functional representations for Mars.

Section 3 reviews some of the radiometric analysis techniques employed in geology. The development of topographic information including crater distribution functions and topographic profiles, the use of global reflectivity in defining geologic units and in establishing surface composition and the interpretation of local reflectivity differences are discussed.

Our analysis on Mariner VI and VII imagery to interpret local variations in reflectivity by using reflectivity ratios is presented in Section 4. We identify the areas where bi-band spectral data is available

and present specific data in support of the interpretation methodology used. The results of our analysis show that light and dark markings interior to crater floors observed in Mariner frames 6N11 and 6N13 acquired through green and red filters respectively are caused by differences between particle sizes between the dark and light areas. Based upon laboratory spectral reflectivity data the most likely size groups for the particles in the two areas is consistent with that proposed by others. Global reflectivity differences between dark and light areas obtained by other authors support larger particle size differences. Based upon our data and that provided by other authors, we concluded that the most likely composition of this local area is limonite stained pyroxene, the latter being a basic silicate common in meteorites. Despite the evidence for basaltic flows on Mars, olivine which is associated with such flows was ruled out as a major constituent in this area based upon its relatively high normal reflectivity.

The application of reflectivity ratio analysis to geologic mapping and the study of atmospheric phenomena was inhibited by the limited amount of data available. It was anticipated that the Mariner IX mission would provide a wealth of bi-band coverage areas on the martian surface and analysis of this data was anticipated. Unfortunately the filter wheel became inoperative on Revolution 118 subsequent to the subsiding of the global martian dust storm. As a consequence, no significant bi-band coverage area of the martian surface was obtained during the Mariner IX mission and this portion of the study was deleted.

## SECTION 2

### THE RADIOMETRIC MODEL

While the primary objective of this study was not the development of basic radiometric (or photometric) principles or techniques but their application to planetology, an understanding of these principles had to be achieved in order to assess the efforts of related research. Unfortunately, the historical development of radiometric science in astronomy and in electro-optics has not had complete commonality--leading to a diversity of definitions and structure that can be confusing to the engineer not versed in both disciplines. This can result in the misinterpretation of the data presented in other research efforts. Nicodemus (1967) has noted:

"Radiometry and particularly its overshadowing subdivision, photometry, are embarrassed by diversity in nomenclature. Careful attention to the definitions of all radiometric terms, symbols, and units both by authors and readers is needed to avoid confusion and misunderstanding."

In this section of our report we present a proposed radiometric model and provide definitions that include concepts from both astronomical and electro-optical radiometry as a basis for standardization in this and future efforts in terrestrial geology as well as planetology.

An important distinction is made depending on the spectral response of the detector in radiometry. Terms such as luminance, illuminance and brightness are used in photometry where the detector has a spectral passband equal to that of the eye (photopic response). Although this is quite precise the prefix "photo" is frequently used when this is not the case.

An example in astronomy and planetology is the definition of "photometric function" which can be measured over a narrow visible or IR spectral band and is rarely measured with a photopic response. The model and terminology presented below is based upon a summary of radiometry (Nicodemus, 1967).

The geometry of the incident and reflected light from a planar surface (in the x-y plane) is shown in Figure 1. Reflectance of an opaque surface is a function of the influx direction through angles  $\theta_i$  and  $\phi_i$  and the efflux direction through angles  $\theta_r$  and  $\phi_r$ . The basic quantity

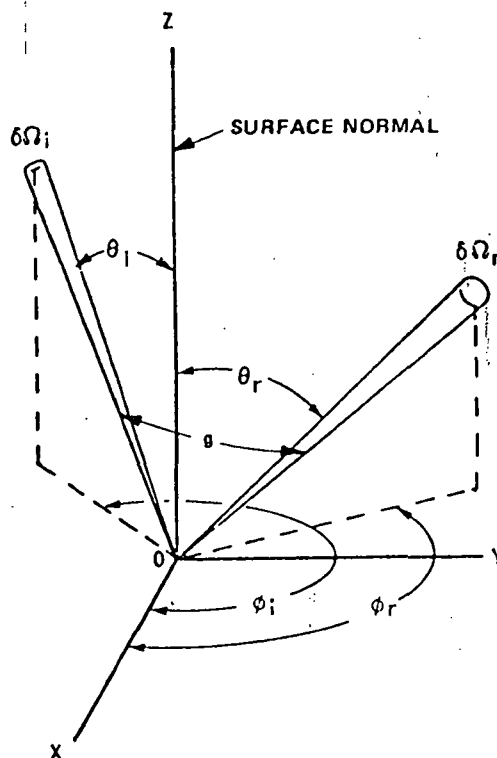


Figure 1. RADIOMETRIC MODEL GEOMETRY

is the spectral radiance,  $N_\lambda(\theta_i, \phi_i, \theta_r, \phi_r)$ , emanating from the surface and is expressed in units  $\text{watts/m}^2 \cdot \text{sr} \cdot \mu\text{m}$ . The wavelength dependence is removed by integrating over the bandwidth of the incident light and the spectral response of the detector. The explicit dependence on wavelength is suppressed below with the understanding that the terms imply a particular bandpass.

The reflectance (dimensionless) of the surface is defined as the ratio of the reflected radiant power to the incident radiant power

$$\rho \equiv \frac{P_r}{P_i} = \frac{\int N_r(\theta_r, \phi_r) d\Omega'_r}{\int N_i(\theta_i, \phi_i) d\Omega'_i} = \frac{\int N_r(\theta_r, \phi_r) d\Omega'_r}{H_i(\theta_i, \phi_i)} \quad (1)$$

where the projected solid angled  $d\Omega' = \sin\theta \cos\theta d\theta d\phi$ . Both the incident radiance  $N_i$  and the reflected radiance  $N_r$  are functions of their respective direction angles  $(\theta, \phi)$ . The reflectance of the surface as defined above can change as the receiver geometry is changed. Clearly we do not generally want to adopt such a definition. An exception occurs when the detector measures all of the reflected radiation, i.e., that reflected into the hemisphere, in which case we obtain the diffuse reflectivity or the Bond albedo discussed below.\*

The reflecting properties of the surface are more appropriately described by the bidirectional reflectance function defined by

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\*The term reflectivity implies decimal fraction while reflectance is expressed in percent. For historical interest we note that the Bond albedo was introduced in 1861.

$$\rho'(\theta_i, \phi_i, \theta_r, \phi_r) \equiv \frac{N_r(\theta_r, \phi_r)}{H_i(\theta_i, \phi_i)} \quad (2)$$

where  $H_i(\theta_i, \phi_i)$  is the incident irradiance in watts/m<sup>2</sup>. Note that  $\rho'$  has dimensions of steradians<sup>-1</sup>. Combining equations (1) and (2) the diffuse reflectivity (or Bond albedo) is given by:

$$\rho_d = \frac{\int_h \rho'(\theta_i, \phi_i, \theta_r, \phi_r) H_i(\theta_i, \phi_i) d\Omega'_r}{H_i(\theta_i, \phi_i)} = \int_h \rho'(\theta_i, \phi_i, \theta_r, \phi_r) d\Omega'_r \quad (3)$$

where "h" signifies integration over the hemisphere. For a perfectly diffuse (Lambertian) surface  $\rho'$  is independent of receiver coordinates and

$$\rho_d = \rho' \int_h d\Omega'_r = \pi \rho' \quad (4)$$

The normal reflectivity (or normal albedo),  $\rho_o$ , is the value of the bidirectional reflectance function when  $\theta_i = \theta_r = 0$  and the influx and efflux directions are both normal to the surface:

$$\rho_o = \frac{N_r(0,0)}{H_i(0,0)} \quad (5)$$

In practice the normal albedo can be computed by measuring the radiance of the center of the planet (or the sub-earth point) near opposition and calculating  $H_i$ .

If we normalize the bidirectional reflectance to 1.0 at zero influx and efflux zenith angles Equation (2) becomes:



$$\rho'(\theta_i, \phi_i, \theta_r, \phi_r) = \rho_o \Phi(\theta_i, \phi_i, \theta_r, \phi_r) \quad (6)$$

where  $\Phi$  is the normalized function. In planetology and astronomy  $\Phi$  is referred to as the photometric function even though it can be measured over non-photopic spectral bandpasses. In radiometric terms it is the normalized bidirectional reflectivity. The diffuse reflectivity (Bond albedo) can now be expressed in terms of the normal albedo, namely

$$\rho_d = \rho_o \int_h \Phi(\theta_i, \phi_i, \theta_r, \phi_r) d\Omega_r' \quad (7)$$

where the integral is referred to as the phase integral.

Retroreflectivity (or geometric albedo) is the value of the bidirectional reflectivity when the influx and efflux angles are equal (zero phase angle), but not necessarily zero themselves, namely

$$\rho_r(\theta) = \rho_o \Phi(\theta, \phi, \theta, \phi) \quad (8)$$

Note that the normal reflectivity is a special case of the retroreflectivity.

The geometric albedo accounts for the variation in apparent brightness across a planetary disk near opposition. Because the lunar disk appears to be uniform, i.e. no limb darkening, its retroreflectivity is constant and equal to its normal reflectivity. This is not true for the martian disk where limb darkening effects have been observed. Frequently the terms normal albedo and geometric albedo are used interchangeably leading to some confusion. The average reflectivity of a planetary disk near

opposition is an area weighted average of its retroreflectivity or geometric albedo but is sometimes identified as the geometric albedo of the planet.

In many situations the bidirectional reflectivity, Eq. (2), is not a function of the four angles  $(\theta_i, \phi_i, \theta_r, \phi_r)$  but only three  $(\theta_i, \theta_r, g)$  where  $g$  is the phase angle shown in Figure 1.\* This allows the bidirectional reflectivity to be written as:

$$\rho'(\theta_i, \theta_r, g) = \frac{N_r(\theta_r, g)}{H_i(\theta_i)} = \rho_r \Phi(\theta_i, \theta_r, g) \quad (9)$$

This simplification is a consequence of assumed symmetry properties of the material, e.g. that the reflectivity is independent of a rotation about its normal. Although most radiometric ("photometric") studies of planetary surfaces accept this assumption it may not always be valid--for example if wind blown dust were to have a preferential deposition  $\Phi(\theta_i, \theta_r, g)$  could be ambiguous.

Finally using the more conventional notation,  $\theta_i = i$  and  $\theta_r = \epsilon$ , the component of radiance due to a planetary surface is given by the expression:

$$N(i, \epsilon, g, \lambda) = S(i, \lambda) t(\epsilon, \lambda) \rho_c(\lambda) \Phi(i, \epsilon, g, \lambda) \quad (10)$$

where  $S(i, \lambda)$  is the irradiance on the surface at wavelength  $\lambda$  and  $t(\epsilon, \lambda)$  is the

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\*Note that  $|\theta_i - \theta_r| \leq g \leq \theta_i + \theta_r$

loss factor due to the transmissivity of the optical path. Note that in the presence of a significant atmosphere  $S$  will include both solar and sky components and that the apparent radiance of the planet will include an atmospheric component as well as a surficial component. For the moon, Mercury and Mars (except during dust storms) the surficial component dominates the observed radiance.

Unfortunately one of the other ambiguities in the definition of radiometric terms is that factors of  $\pi$  "come-and-go". Some authors may have a factor of  $1/\pi$  multiplying the right side of Eq. (10). This factor originates because of a difference in the definition of photometric units between the English and metric systems. To further complicate the situation, a similar difference does not formally exist for radiometric units although many authors impose such conditions in order to use operationally measured reflectivities. The English units were developed to facilitate the operational measurement of reflectivity such that a perfectly reflecting diffuse (Lambertian) surface has a bidirection reflectivity of unity. The English unit of luminance (ft-Lambert) has been rationalized by a factor of  $\pi$  in addition to the conversion factor from  $m^2$  to  $ft^2$ . Consequently the luminous bidirectional reflectivity, Eq. (2), is increased by  $\pi$ , viz:  $\rho' = \pi N_R/H_i$  and the diffuse reflectivity, Eq. (4), becomes  $\rho_d = \rho'$  (i.e. the bidirectional reflectivity of a Lambertian surface is numerically equal to its diffuse reflectivity) and the factor of  $1/\pi$  enters Eq. (10). Note also that  $N_r(\theta_r, \phi_r)$  is constant for a Lambertian surface and the definition based upon the ratio of reflected to incident power, Eq. (1) also becomes  $\rho = \pi N_r/H_i$  and the two definitions of luminous reflectivity are identi-

cal for Lambertian surfaces. Operationally we frequently measure the bidirectional reflectivity of a material by comparing the received power to that received from a Lambertian standard (with  $\rho_d \approx 1.0$ ) at the same orientation, the value obtained is governed by the alternate definition discussed here and the multiplying factors discussed above apply.

### SECTION 3

#### BIDIRECTIONAL REFLECTIVITY (PHOTOMETRIC) FUNCTION

It is clear from the results of the previous section that some knowledge of the bidirectional reflectivity characteristics of the surficial materials on a planet are required in order to interpret the radiance measured by either an imaging system or a radiometer. Even if a planetary surface were composed of a single type of material with a well defined particle size regime, the radiance of the planetary surface will vary due to geometrical considerations alone. Both the normal reflectivity and the normalized bidirectional reflectivity (photometric) function can change due to compositional differences on the planetary surface. Consequently, it becomes a matter of removing the geometrical dependence of the radiance to arrive at an "albedo" which contains information about the composition of the surface. One approach is to use Eq. (10) to convert the observed radiance into the normal albedo. This requires knowledge of the normalized bidirectional reflectivity or photometric function. The normal albedo differences subsequently derived can be due to chemical or mineralogical differences, particle size differences or a combination of these effects. A primary objective of this study was to determine whether observed albedo differences (dark and light markings) in Mariner VI and VII photography of the Martian surface were due to chemical effects or particle size differences. However the analyses methodology did not require the computation of the normal albedo or reflectivity.

If we assume that there are no chemical or composition differences then the photometric function or normalized bidirectional reflectivity can be derived from measurements of the radiance of the surface at different illumination and viewing conditions. The measured data is fit to an analytic model for the photometric function. The most frequently used functional representation for surface radiance is the Minnaert equation (Minnaert, 1941). This equation is a parametric expression developed to obey the Helmholtz reciprocity law -- the bidirectional reflectivity is invariant upon reversability of the incident and emission angles, namely

$$\rho'(i, e, g) = \rho'(e, i, g) \quad (11)$$

The resulting parametric equation has the form

$$N(i, e, g, \lambda) = N_p(g, \lambda) (\cos i)^{k(g, \lambda)} (\cos e)^{k(g, \lambda) - 1} \quad (12)$$

where the exponential parameter,  $k$ , is a function of both the phase angle and wavelength and can be thought of as a limb darkening parameter. The moon, which has no limb darkening effect, has a value of  $k=0.5$  at zero phase while a Lambertian surface would have a value of  $k=1.0$ . The Minnaert law is frequently written with the symbol "B" in place of the symbol "N" which we have employed. The former implies the brightness of the surface in contrast to the radiance and consequently in keeping with the discussion in Section 3 we have avoided its use here. The term  $N_p(g, \lambda)$ , frequently written as " $B_0$ " can be a function of phase angle without violating the reciprocity requirement

since it does not depend on the incident or emission angle. At zero phase note that  $N_p$  is equal to the radiance obtained at normal incidence and normal emission, namely  $N_p(0,\lambda) = N_r(0,0) = H_o(\lambda) \rho_c$  where  $H_o$  is the irradiance at normal incidence. This is required so that Eq. (12) has the correct limit at  $i = \epsilon = 0$ . The Minnaert equation is frequently used in a modified form obtained by implicitly dividing both sides by  $H_o(\lambda)$ . In this case the value of the parameter  $N_p$  or  $B_o$  at zero phase angle is the normal reflectivity or albedo. If we assume that  $N_p$  is not a function of the phase angle, then  $N_p = H_o(\lambda) \rho_c$  for the reasons given above. Some authors make this assumption without noting that it may not be justified.

Comparing Eq. (12) to Eq. (10) and assuming that the loss due to the atmosphere is negligible, the photometric function described by the Minnaert equation is

$$\bar{\Phi}(i, \epsilon, \xi, \lambda) = \frac{N_p(\xi, \lambda)}{H_o(\lambda) \rho_c(\lambda)} (\cos i)^{k(\xi, \lambda)-1} (\cos \epsilon)^{k(\xi, \lambda)-1} \quad (13a)$$

Equating the photometric function and the Minnaert equation, the practice of some authors, is somewhat imprecise and misleading. If  $N_p$  is not a function of the phase angle the photometric function becomes

$$\bar{\Phi}(i, \epsilon, \xi, \lambda) = (\cos i)^{k(\xi, \lambda)-1} (\cos \epsilon)^{k(\xi, \lambda)-1} \quad (13b)$$

The Far Encounter photography obtained during the Mariner VI and VII missions has been employed to estimate the parameters  $N_p, k$  in Equations (12) and (13) (Young, 1971). The video signal was converted to reflectivity using available calibration data such that the value of  $N_p/H_0$  at zero phase would be the normal reflectivity (albedo),  $\rho_0$ , if it could have been estimated. Log-log plots of  $N \cos \epsilon$  versus  $\cos i \cos \epsilon$  called "Minnaert plots" were constructed to estimate  $N_p/H_0$  and  $k$ . Since these photographs were obtained with the long focal length TV camera and occurred at limited phase angles,  $\sim 23^\circ \pm 2^\circ$  this evaluation did not include the variation of  $N_p/H_0$  or  $k$  with phase angle or wavelength. The results obtained from each of the missions are summarized in the table below. The variation of  $k$  from place to place on Mars was attributed to a variation in composition and supported the hypothesis that the Martian surface is more than a two-component system.

Table 1  
Minnaert Parameters from Mariner VI and VII Imagery

Region	West Longitude, deg.	Latitude, deg.	$k_6$	$k_7$	$\frac{N_{p6}}{H_0}$ Ster <sup>-1</sup>	$\frac{N_{p7}}{H_0}$ Ster <sup>-1</sup>
Ophir	68	-8	0.63	0.71	0.146	0.131
Center of Elysium	213	+12	0.56	0.55	0.144	0.132
Aeolis	213	0	0.61	0.68	0.133	0.133
Center of Syrtis Major	290	+5	0.46	0.48	0.093	0.071
Solis Lacus	90	-40	0.66	0.60	0.117	0.093



Earth-based observations of the radiance of Mars have also been employed to estimate the parameter  $k$  (Binder, 1972). Unfortunately Earth-based observations are constrained in that measurements cannot be made of the Martian surface at phase angles greater than  $40^\circ$  due to ephemeris considerations. The results of these measurements show a slightly higher value of  $k$  at a  $23^\circ$  phase angle, namely  $k \approx 0.8$ . These data were measured at a wavelength of  $0.60 \mu\text{m}$ . The differences between these two measurements of  $k$  are most likely the result of the differences in the spectral bandpass of the sensors. This study also evaluated the variation of  $k$  with wavelength with the resulting dependence at approximately  $10^\circ$  phase angle.

$$k(10^\circ, \lambda) = 0.91 - .16/\lambda \quad (14)$$

Since the spectral passband of the B camera varies from  $0.48 \mu\text{m}$  to approximately  $0.65 \mu\text{m}$  (Danielson, 1971) the value of the  $k$  parameter according to Eq. (14) corrected for the phase angle difference would vary from 0.74 to 0.84, still somewhat higher than the Mariner derived values. We note, however, that Binder arbitrarily assumed that  $k$  at zero phase angle was 0.5 (no limb darkening effect).

A more recent evaluation of the Minnaert coefficients derived from observations on the Mariner IX imagery (Thorpe, 1973) resulted in values presented in Table 2 as a function of phase angle at an effective wavelength of  $0.56 \mu\text{m}$ . In this case also the value of  $N_p$  at zero phase would be the normal reflectivity. Note that the results presented in the table clearly indicate that  $N_p$  is dependent on the phase angle. If we fit results of

Table 2  
Minnaert Parameters from Mariner IX Imagery

Phase (deg)	20	30	40	50	60	70	80
$N_g$	$0.166 \pm 0.02$	$0.156 \pm 0.30$	$0.145 \pm 0.03$	$0.14 \pm 0.02$	$0.132 \pm 0.03$	$0.129 \pm 0.03$	$0.135 \pm 0.04$
$k$	$0.70 \pm 0.01$	$0.73 \pm 0.06$	$0.77 \pm 0.03$	$0.80 \pm 0.02$	$0.84 \pm 0.05$	$0.87 \pm 0.04$	$0.91 \pm 0.05$

Thorpe for the variation of  $k$  with phase angle with a linear equation we find that

$$k(g) = 0.63 + 0.0035 g \quad (15)$$

since  $k(0) > 0.5$  this indicates a limb darkening effect on Mars.

Measurements of the Minnaert coefficient  $k$  at a mean phase angle of  $58^\circ$  were also made during the initial portion of the Mariner IX mission while a dust storm covered Mars (Masursky et al, 1972). For these measurements  $k$  increased from 0.93 to 1.12 with increasing wavelength and indicate a nearly Lambertian surface ( $k=1.0$ ). This result would be expected for an optically thick atmosphere composed of small particles ("dust").

A recent research effort has led to an alternative expression for the photometric function (Meador, 1975). Although more complex than the Minnaert law the Meador equation has the distinct advantage that its parameters can be related to the physical characteristics of the reflecting surface, e.g. particle size, single-particle albedo and compactness. The explicit functional form will not be presented here and the interested reader is referred to the cited reference for details. This expression has been compared to the three data sets previously used to derive the parameters for the Minnaert law as discussed above (Weaver, 1974). Based upon the fit of the Meador expression to the experimental observations and laboratory data on Colorado Basalt it was concluded that the mean intercenter spacing

of adjacent particles is about 4/3 of the mean diameter. The mean diameter was concluded to be greater than  $225 \mu\text{m}$  and not in conflict with other indications that the mean diameter of Martian surface particles is about  $400 \mu\text{m}$ .

Regardless of which expression is used to represent the bidirectional reflectivity or photometric function in order to analyze the radiance of a planetary surface as measured by either an imaging system or a radiometer, the values of the influx or incident angle,  $i$ , the efflux or emission angle,  $\epsilon$ , and the phase angle,  $g$ , must be determined. If the latitude and longitude of the sub-solar point are known  $(\phi_{s_z}, \lambda_{s_z})$  the incident angle at a point having latitude and longitude  $(\phi, \lambda)$  is given by

$$\cos i = \sin \phi \sin \phi_{s_z} + \cos \phi \cos \phi_{s_z} \cos(\lambda - \lambda_{s_z}) \quad (16)$$

The sub-solar point is the location on the planetary surface where the line from the center of the sun to the center of the planet intersects the surface. Similarly if the sub-spacecraft point is known  $(\phi_{ss}, \lambda_{ss})$  the emission angle can be determined by using

$$\cos \epsilon = \sin \phi \sin \phi_{ss} + \cos \phi \cos \phi_{ss} \cos(\lambda - \lambda_{ss}) \quad (17a)$$

As the spacecraft approaches the planet a (parallax) correction may be required. The emission angle,  $\epsilon$ , given by Eq. (17a) is increased by  $\delta$ .

$$\tan \delta = \sin \epsilon / (R_s/r - \cos \epsilon) \quad (17b)$$

where  $\epsilon$  is the value determined from Eq. (17a),  $R_s$  is the distance from the spacecraft to the center of the planet and  $r$  is the radius of the planet. For the Mariner VI and VII imagery the values of  $R_s$  have been computed by Davis (1971) and  $r = 3385\text{km}$  for Mars. For the Mariner VI and VII imagery  $R_s/r = 2-3$  and the correction factor is significant. The locations of the sub-solar and sub-spacecraft points are not as readily documented and must be determined from orbital data. If the azimuth angle between the incident and emission planes is known and equal to  $\theta$  then the phase angle is given by

$$\cos \phi = \cos i \cos \epsilon + \sin i \sin \epsilon \cos \theta \quad (18)$$

The angle  $\theta$  is the difference between the bearing of the sub-solar point,  $Z_{sz}$ ,

$$\cos |Z_{sz}| = \frac{\sin \phi_{sz} - \cos i \sin \phi}{\sin i \cos \phi} \quad (19a)$$

and the bearing of the sub-spacecraft point,  $Z_{ss}$ ,

$$\cos |Z_{ss}| = \frac{\sin \phi_{ss} - \cos \epsilon \sin \phi}{\sin \epsilon \cos \phi} \quad (19b)$$

The sign of  $Z_{sz}$  and  $Z_{ss}$  is determined by the value of the longitudes  $\lambda_{sz}, \lambda_{ss}$  with respect to  $\lambda$ .

## SECTION 4

### RADIOMETRIC APPLICATIONS IN GEOLOGY

In this section we discuss some of the applications of radiometry and the bidirectional reflectivity function in obtaining information about the geology of planets. No attempt has been made to make this discussion exhaustive but only representative of some of the applications available. The discussion has been divided into three categories which include (1) the influence of reflectivity on topographic information, (2) the use of reflectivity for global geology, and (3) the use of local changes in reflectivity in assessing compositional differences in surficial material.

#### 4.1 Topographic Applications

Two examples where the reflectivity characteristics of a planetary surface influences topographic information include the use of imagery in crater distribution studies and the development of topographic profiles from photometrically corrected image data.

The effect of the finite resolution of an imaging sensor (e.g., Mariner television systems) upon the accuracy of distribution data is well known, namely as the diameter of the crater approaches the resolution limit of the imaging system the measured distribution falls below the actual distribution due to the inability to identify and measure the craters on the reconstructed imagery. The reflectivity characteristics of the planetary surface play a role in how rapidly the loss of information due to limited

resolution occurs. This point is not necessarily appreciated by all investigators utilizing the imagery in obtaining their photogrammetric information. The resolution of an imaging system is most often quoted assuming that the object being imaged is high contrast (having extreme radiance values). As the contrast of the object decreases the resolution limit becomes poorer. That is, the inherent resolution limit of the imaging system is only achieved under illumination and viewing conditions where the objects under study have a reasonably high contrast. At high phase angles most surface features have associated shadows which enhance their contrast. Consequently, one would expect the information content of imagery acquired at large phase angles to be higher than that acquired at lower phase angles. As part of the development of quality evaluation techniques for Lunar Orbiter photography the author (Kinzly, 1967) developed an expression for information content in terms of a number of mission parameters including phase angle which was subsequently related to the minimum detectable crater diameter of Lunar Orbiter photography. The variation of the minimum crater diameter with respect to phase angle for this photography is shown in Figure 2. A 4:1 variation occurs as the phase angle changes from 0 to  $90^{\circ}$ . The change is most dramatic in the region from  $60^{\circ}$  to  $90^{\circ}$  emphasizing the value of large phase angle photography. More recently the importance of illumination conditions upon comparative planetary geology has been evaluated (Schultz, 1976). Comparative geology uses imagery from several planetary and lunar missions. It is noted, for example, that the information content of the Mariner IX imagery is comparable to the Mariner VI and VII imagery despite the smaller slant ranges involved on the former mission. The improved theoretical

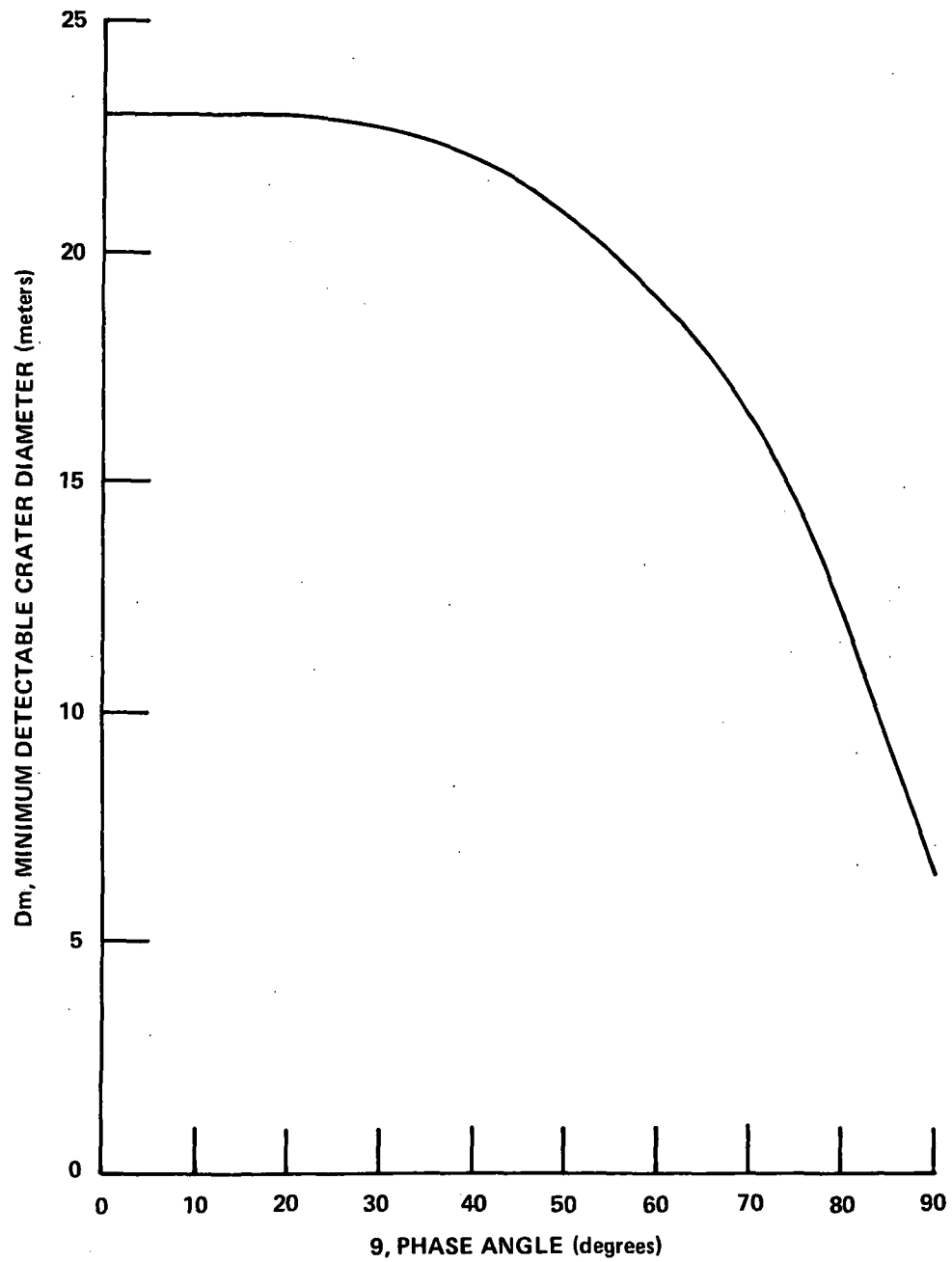


Figure 2 EFFECT OF PHASE ANGLE ON INFORMATION CONTENT OF NEAR VERTICAL LUNAR ORBITER PHOTOGRAPHY



resolution limit of the Mariner IX imagery is offset by the smaller phase angles at which the imagery was obtained. A comparison of TV and photographic imagery of the moon shows that image enhancement allows the TV image to approach the limiting resolution of 2.2 TV lines. However, the enhancement can distort topographic features which are on the order of several resolution elements.

Since the radiance of a point on a planetary surface depends upon the incident and emission angles measured with respect to the local surface normal, an accurate measure of surface radiance can be interpreted in terms of the surface orientation if the bidirectional reflectivity function for the surface is known. The conversion of surface radiance measured from an image into local slope and subsequently into a topographic profile has been termed "photoclinometry" and was applied to both Ranger (Rindfleisch, 1966) and Lunar Orbiter (Lambiotte, 1967) imagery. In the Lunar Orbiter application, the technique was used to derive relative surface roughness indices at potential Apollo landing sites. An evaluation of the quality of Lunar Orbiter imagery (Kinzy, 1968) showed that the medium resolution photographs of Missions I and II received excessive exposure and the image densities could not be reliably converted to surface radiance. The limitation that this placed on the application of photoclinometry was investigated (Gambell, 1968). More recently photoclinometry was used on Mariner IV imagery to produce depth/diameter data for martian craters (Cintala, 1976). Because of the nonlinear relationship between surface radiance and surface slope, the average radiance measured by an imaging system is not necessarily equivalent to the average slope of the surface -- both averaged over a resolution element. Consequently,

the resolution of the imaging system can affect the accuracy of the profile information generated by photoclinometry. This analytical procedure, to the best of our knowledge, has never received widespread application beyond Lunar Orbiter.

#### 4.2 Definition of Global Geology

As noted in Section 3, differences in the normal albedo of a planetary surface can be produced by chemical or mineralogical differences, particle size differences or a combination of these effects. Consequently, a geologist frequently makes use of the normal albedo of the surface as an aid in defining geologic units. This potential utilization led to the development of albedo maps of the moon (Pohn, 1970), Mars (Cutts, 1971 and de Vaucouleurs, 1973) and Mercury (Dzurisin, 1976). In addition to the fact that albedo boundaries can be coincident with the boundaries between different geological units, correlation between reflectivity of surficial materials has been noted to vary inversely with age on the lunar surface -- older and more heavily cratered areas having a higher reflectivity.

In addition to the production of albedo maps which are used in qualitative geologic evaluations, analysis of the reflectivity characteristics of the light and dark areas of a planet on the global scale have been used to draw inferences on the composition of surficial materials. Typical examples of these type of analyses have included those of Pollack and Sagan (1967) where it was concluded that the bright and dark areas of Mars appeared to have a very similar chemical composition with goethite, a major constituent of both

areas. It was also suggested that both areas were covered by a fine powder and that the seasonal change in the reflectivity of the dark areas was produced by aeolian transport of material to and from the dark areas. The average diameter of the particles in the bright areas is estimated to be  $50\ \mu\text{m}$  while the dark areas were estimated to have particles ranging from  $200\ \mu\text{m}$  to  $400\ \mu\text{m}$  depending upon the period relative to the seasonal darkening. Based on visible polarimetry and IR radiometry they proposed limonite as a primary constituent of this powder.

A second example of global reflectivity analysis is that of Binder and Jones (1972). Observations were made using a ten channel spectroradiometer operating from  $0.60 - 2.27\ \mu\text{m}$ . They concluded that the measured spectro-reflectivity fell into two well-defined groups, mare and desert units, indicating that the martian surface consists of two types of materials. By comparison with laboratory data they proposed that both surface units contained lithic soils having a limonite stain and that the soils of the maria are richer in pyroxene, olivine, or both than the desert soils. This investigation, therefore, suggests that the differences between martian maria and deserts are due to composition and not due to differences in particle size as proposed by Pollack and Sagan.

A third example of global reflectivity investigations is the work of McCord and his associates (McCord and Adams, 1968; Adams and McCord, 1969; McCord and Westphal, 1971). These investigations involved spectroradiometric observation of Mars from  $0.30 - 2.5\ \mu\text{m}$ . Comparison of the resulting spectro-

reflectivity curves to laboratory data led to the conclusion that oxidized basalts are a significant constituent of the martian surface. Evidence has been suggested for the existence of basaltic flows and mare on Mars, thereby supporting this hypothesis. It should be noted, however, that the Mariner IX imagery acquired after a martian dust storm contained numerous bright and dark areas produced by dynamic aeolian transport of material over the martian surface. These observations support the hypothesis by Pollack and Sagan that the seasonal changes in reflectivity are a result of a change in the average particle size. It is, of course, quite possible that both mechanisms play a role in the dynamic characteristics of the reflectivity of the martian surface.

#### 4.3 Interpretation of Local Reflectivity Differences

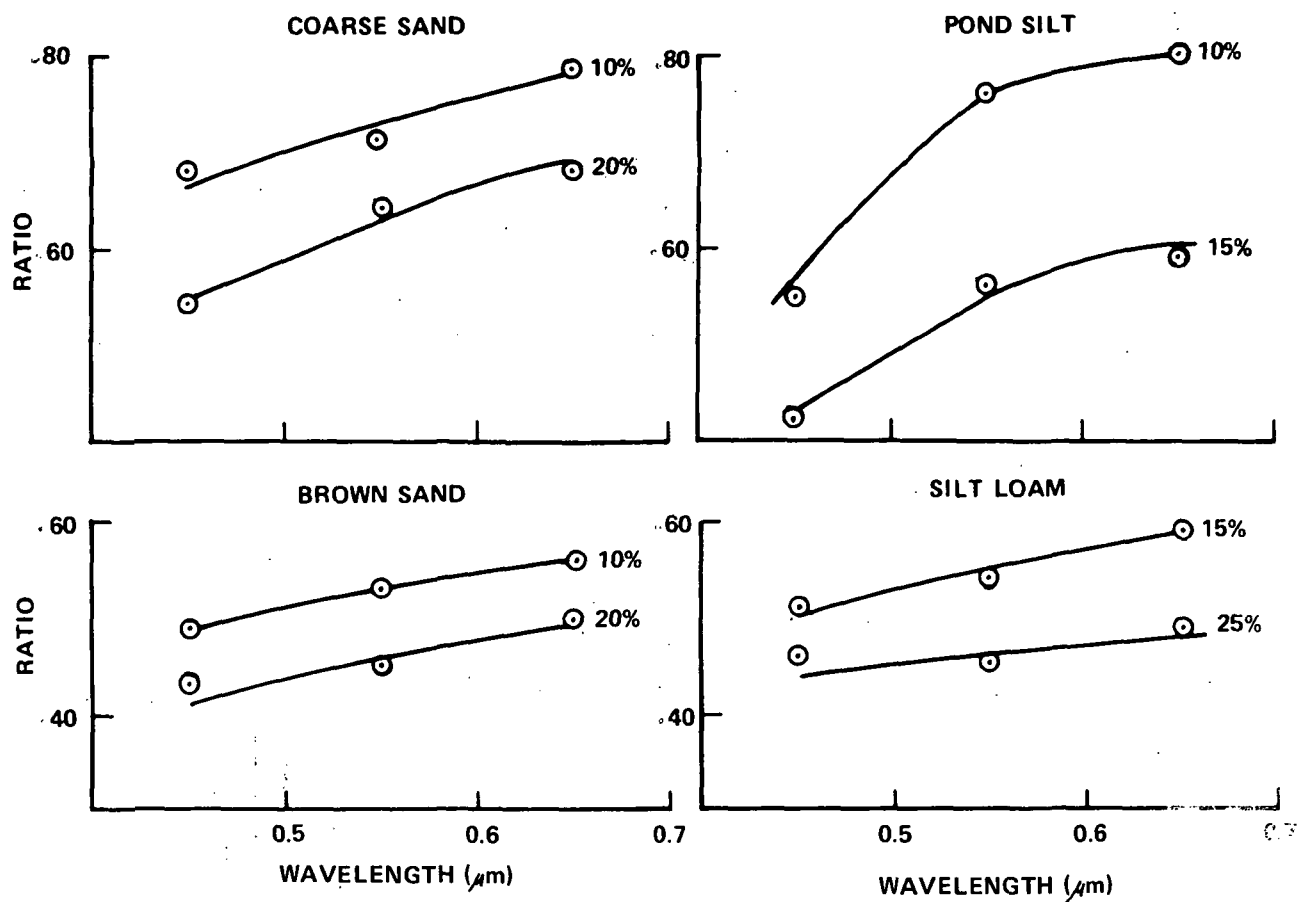
The evaluation of local changes in planetary surface reflectivity requires the availability of radiometrically calibrated high resolution imagery. Consequently, the techniques available for such analyses are less developed although they parallel those used in evaluation of reflectivity changes on a global scale. The distinction that we make between global and localized analyses can be correlated directly with earth-based versus spacecraft observations of the planet. An example of the qualitative analyses of local reflectivity variations is the use of the radial bright ray patterns associated with some craters on the lunar surface in order to derive relative ages. Quantitative techniques are available as illustrated in this report to analyze the dark and light markings occurring in the Mariner VI and VII imagery.

As noted in the previous section, Young and Collins (1971) analyzed the Mariner VI and VII far encounter photographs to derive the parameters of the Minnaert photometric function for five regions of the martian surface. We consider this analysis to be global rather than local, however, since it only employed the far encounter pictures which have resolutions ranging from 10 - 100 km.

A method for evaluation of local changes in the reflectivity of soils employing radiometric analysis using conventional color aerial photography has been recently reported (Piech, 1974). The technique was developed to assist in terrestrial soil surveys by supplementing conventional land form analyses with reflectivity information extracted from color imagery. The cause of reflectivity variations in a soil unit are evaluated by computing the ratio of the reflectance in two spectral bands. Information on the relative soil moisture and texture characteristics is derived from the reflectivity ratios obtained from the imagery. In this case the aerial camera is employed as a radiometer. In addition to calibration, of the object-to-image radiometric response, the effects of the atmosphere must be removed from the imagery in the earth-based application. This is accomplished by radiometric analysis of standard scene objects such as shadow areas. If the darker soil element has a greater red-to-blue reflectivity ratio than the lighter soil element, the reflectivity variation is caused by differences in moisture content. If the darker soil element has a smaller ratio, the decrease in reflectivity is produced by an increase in the average particle size. Rather than computing the reflectivity ratios in two spectral bands the equivalent procedure of

computing the ratio of the reflectivity of the darker unit to the lighter unit in each of the spectral bands selected can be used. If the ratio of the darker to lighter soil element at the longer wavelength (e.g. red) is less than the ratio at a shorter wavelength (e.g. blue) then the change in reflectivity can be attributed to a larger particle size for the darker element. Figure 3 shows laboratory data obtained in support of this analysis technique. The figure shows the ratio of the darker to lighter (i.e. wet to dry) soil elements for several terrestrial sands and varying amounts of moisture content. The reflectivity ratios were obtained from densitometer measurements of vertical photography (normal emission) obtained under ambient sunlight such that the incident angle and phase angle varied from  $30^{\circ}$  to  $50^{\circ}$ . Note that the ratio increases with wavelength supporting the interpretation of reflectivity ratios proposed above. In the case of reflectivity changes due to particle size variation, the ratio decreases with increasing wavelength. Additional experimental data is presented in the cited reference. This size-reflectivity change rule is not applicable to all types of minerals that could be constituents of soils. The reflectivity versus size effect has been investigated for likely martian surface materials (Salisbury, 1968) and is discussed further in Section 5.2.

The application of this analysis technique to adjacent light and dark surficial elements on the martian surface can be used to distinguish between a chemical cause for lower reflectivity, namely absorbed moisture or a physical cause due to varying particle size. Note that if the composition



**Figure 3 WET/DRY RATIOS DETERMINED PHOTOGRAPHICALLY FOR VARIOUS TERRESTRIAL SOILS AND MOISTURE CONTENTS BY WEIGHT**

of the light and dark areas is assumed to be similar or if the bidirectional reflectivity is assumed constant, the ratio of the radiance of the dark to the light area is equal to the ratio of their reflectivities and no explicit correction for the bidirectional reflectivity or photometric function is required. This technique was applied to the analysis of adjacent light and dark areas occurring in the Mariner VI and VII imagery -- the results are presented in the next section of this report.



## SECTION 5

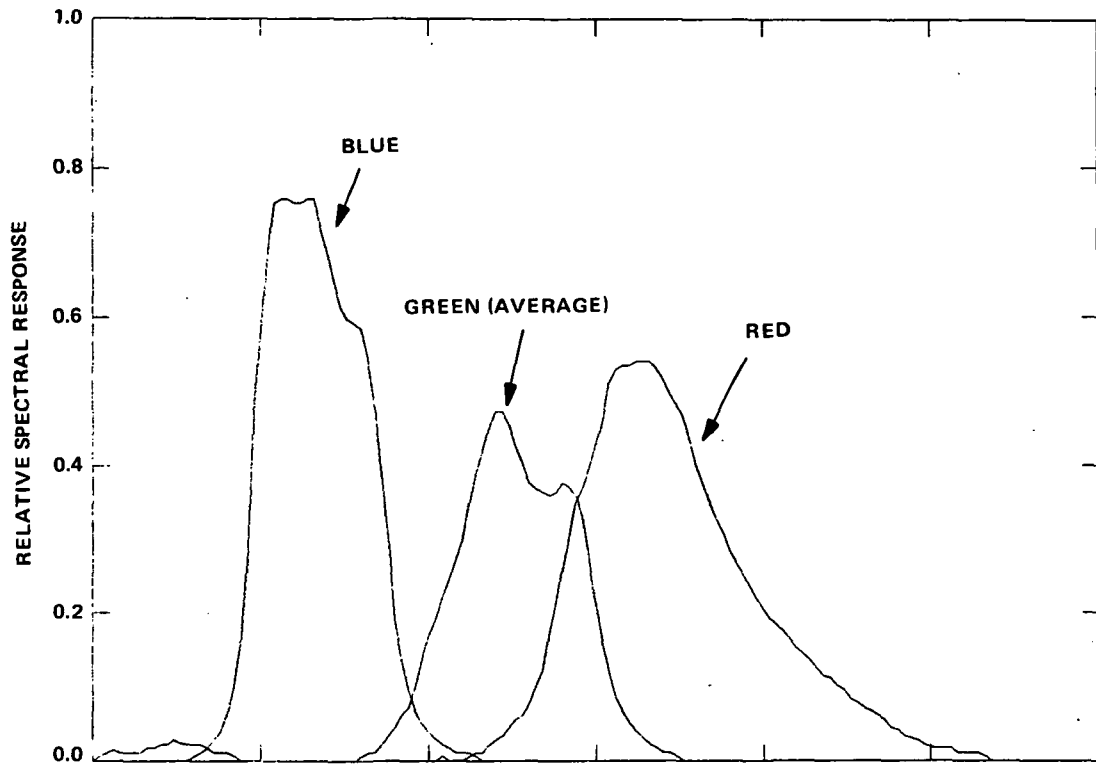
### ANALYSIS OF MARINER VI AND VII IMAGERY

This section of our report documents analyses carried out on Mariner VI and VII imagery to interpret local variations in reflectivity by using reflectivity ratios derived from the imagery. Darkening of the martian surficial material can be produced by chemical effects such as absorbed water, by changes in the particle size or by changes in the composition of the surface material. Identification of the presence of significant amounts of moisture in the martian soil is important to the assessment of life forms on Mars while the assessment of relative particle size differences is important to the evaluation of dynamic aeolian processes. Alternate near encounter frames from the wide angle Mariner cameras were taken through different spectral filters (red, green or blue) and the overlap between successive frames show the same area of the martian surface in two spectral bands. These regions of overlap provide an opportunity to use reflectivity ratio analysis techniques in the study of martian planetology. In this section we (1) identify the areas where bi-band spectral data is available, (2) present additional data in support of the interpretation of the reflectivity ratios beyond that identified previously in Section 4.3, (3) describe the tools developed to compute the reflectivity ratios from digital data supplied by the Image Processing Laboratory at JPL and (4) present the results of the analysis of adjacent light and dark areas occurring in the Mariner photography.

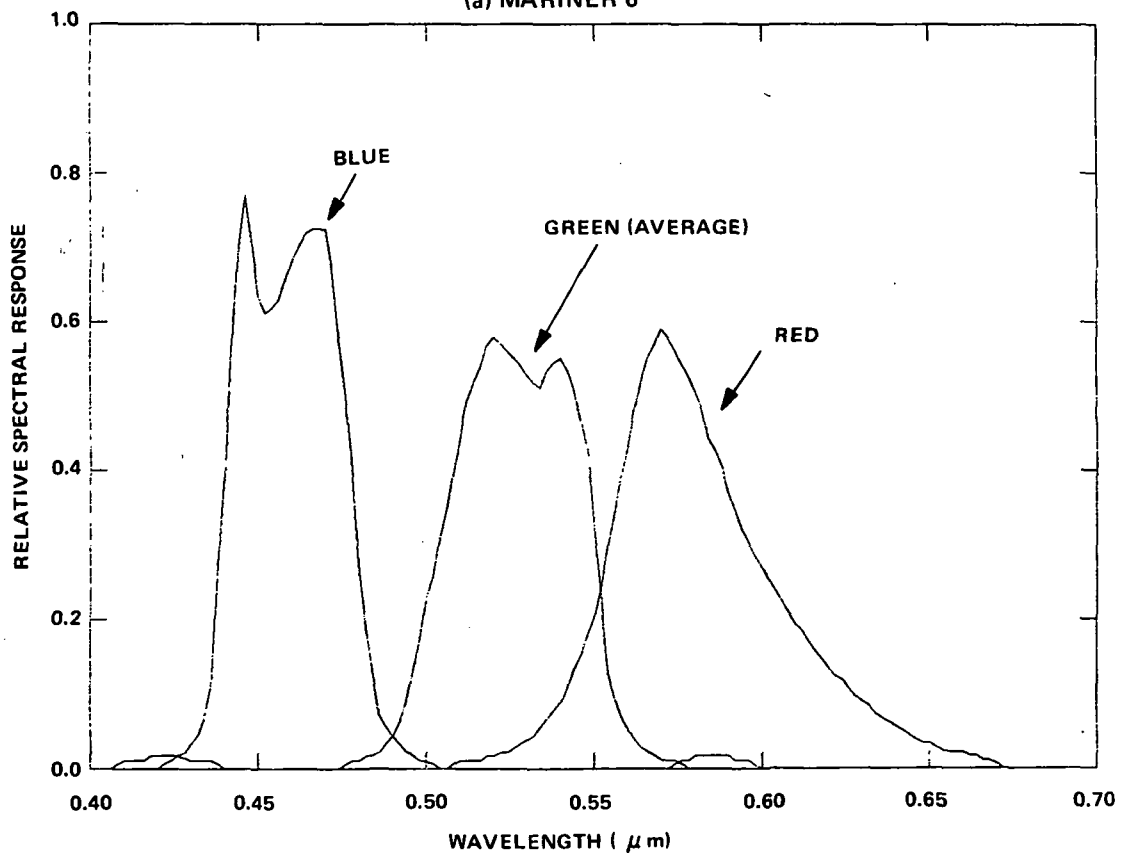
### 5.1 Mariner VI and VII Multi-Spectral Image Data

The objective of this investigation was to illustrate the utility of radiometric analysis, particularly reflectivity ratios, to planetology in general and the study of Mars in particular. We planned to use TV image data obtained from the Mariner VI, VII and IX missions. Mariner VI and VII were fly-by missions that had a variety of experiments including a dual TV imaging system (Danielson, 1971). One TV sensor employed a short focal length lens ( $\sim 50\text{mm}$ ) while the second sensor employed a long focal length lens ( $\sim 500\text{mm}$ ) and were termed the "wide angle" and "narrow angle" cameras respectively. Since the imagery was obtained during fly-bys the resolution ( $\sim 2.2$  TV lines) changed from image to image. For the imagery utilized it was approximately 2-4 km for Mariner VI and 4-5 km for Mariner VII. The wide angle camera contained a filter wheel with four spectral filters. The spectral transmission of the filters incorporated into Mariner VI and VII are shown in Figure 4. Alternate near encounter frames from the wide angle Mariner cameras overlapped such that the same area of the martian surface was taken through two different spectral filters. The initial step in the research effort was the identification of those areas where bi-band coverage was available. Reflectivity ratios computed within these areas could be used for geologic mapping, discrimination of atmospheric phenomena, obtaining information about polar cap structure and obtaining information about the cause of local variations in the reflectivity of the martian surficial material.

Six areas of bi-band coverage were identified from the Mariner VI mission and six areas of bi-band coverage from the Mariner VII mission.



(a) MARINER 6



(b) MARINER 7

Figure 4 SPECTRAL TRANSMISSION OF MARINER VI AND VII WIDE ANGLE CAMERAS

Table 3 summarizes these areas which are designated by the mission number followed by a letter code A through G. This table identifies the location of each of these areas and the frame numbers and spectral filters of each of the overlapping photographs. Figure 5 shows the location of each of the bi-band coverage areas on the 1971 Shaded Relief Map of Mars.

A number of potential bi-band coverage areas were eliminated from consideration. These included the two areas of overlap between 7N5, 7N7 and 7N9. These photographs were obtained at a highly oblique angle and had lower resolution than the subsequent near encounter frames. Consequently the radiometrically corrected digital imagery was not requested from JPL. In retrospect this is somewhat unfortunate since these photographs included some of the areas in Meridiani Sinus coincident with areas 6B and 6C previously selected. In addition, the overlap area between frames 7N29 and 7N31 were deleted from consideration because of their proximity to the terminator which yields a lower quality due to the low exposure level on the vidicon. One area of overlapping coverage among frames 7N15, 7N17 and 7N19 actually had tri-band coverage. This is a small area in the South Polar Cap located at  $72^{\circ}\text{S} - 80^{\circ}\text{S}$ ,  $332^{\circ}\text{W} - 348^{\circ}\text{W}$ . Although no specific analysis was undertaken over this area, several features appear that might be caused by atmospheric phenomena. Figures 6 through 9 show the Mariner VI bi-band coverage areas on enhanced versions of the imagery transformed to orthographic projections. Figures 10 through 13 show the bi-band coverage areas of Mariner VII on similar projections. Digital tapes containing TV image data for each frame identified in Table 3 were obtained from Mr. A. Collins at the

Table 3a  
Mariner B1-band Coverage Areas  
Mariner VI

DESIG- NATION	REGION	LOCATION	FRAME NO.	SPECTRAL BANDS	OVERLAP (APPROX.)	PHASE ANGLE	GEOLOGIC UNIT*	PRINCIPAL MAPPING QUADRANGLE (S)
6A	Aurorae Sinus - Margaritifer Sinus	20°N - 19°S, 28°W - 47°W	6N5, 6N7	Red, Green 1	40%	51°	cu, h	Margaritifer Sinus
6B	Thymiamata	11°N - 13°S, 2°W - 15°W	6N9, 6N11	Blue, Green 2	40%	51°	cu	Oxia Palus - Margar- itifer Sinus
6C	North Meridiani Sinus & West Edom	7°N - 10°S 352°W - 30°W	6N11, 6N13	Green 2, Red	30%	51°	cu	Sabaeus Sinus
6D	Margaritifer Sinus and East	9°S - 23°S, 13°W - 23°W	6N15, 6N17	Green 1, Blue	25%	80°	cu	Margaritifer Sinus
6E	Margaritifer Sinus and East	13°S - 21°S, 1°W - 6°W	6N17, 6N19	Blue, Green 2	25%	80°	cu	Margaritifer Sinus
6F	Sabaeus Sinus and West	12°S - 21°S, 350°W - 353°W	6N19, 6N21	Green 2, Red	25%	80°	cu	Sabaeus Sinus
6G	Sabaeus Sinus and West	11°S - 20°S 338°W - 341°W	6N21, 6N23	Red, Green 1	25%	80°	cu	Sabaeus Sinus

\*List of Units (after Carr, 1973)

h: Chaotic Deposits  
cu: Cratered Deposits, Undivided

Table 3b  
Mariner BI-band Coverage Areas  
Mariner VII

DESIG- NATION	REGION	LOCATION	FRAME NO.	SPECTRAL BANDS	OVERLAP (APPROX.)	PHASE ANGLE	GEOLOGIC UNIT*	PRINCIPAL MAPPING QUADRANGLE (S)
7A	Edge Polar Cap	50°S - 70°S, 2°W - 56°W	7N11, 7N13	Green 2, Red	50°	35°	cm, cu, ps	Argyre
7B	Polar Cap	57°S - 79°S, 342°W - 56°W	7N13, 7N15	Green 2, Red	50°	35°	cm, cu	Argyre - Mare Australe
7C	Polar Cap	64°S - 86°S 322°W - 26°W	7N15, 7N17	Red, Green 1	50°	35°	?	Mare Australe
7D	Polar Cap	68°S - 88°S 292°W - 4°W	7N17, 7N19	Green 1, Blue	50°	35°	?	Mare Australe
7E	Noachis - Hellas	29°S - 44°S 322°W - 342°W	7N23, 7N25	Green 1, Blue	25%	80°	cu	Noachis
7F	Boundary Helle- spontus Hellas Boundary	38°S - 48°S 311°W - 322°W	7N25, 7N27	Blue, Green 2	15%	80°	cu, m, ps	Noachis - Hellas
7G	Hellas Floor	40°S - 45°S 295°W - 303°W	7N27, 7N29	Green 2, Red	small	80°	ps	Hellas

\*List of Units (after Carr, 1973)

cm: Cratered Deposits, Mantled  
cu: Cratered Deposits, Undivided  
ps: Sparsely Cratered Plains  
m: Mountainous Deposits

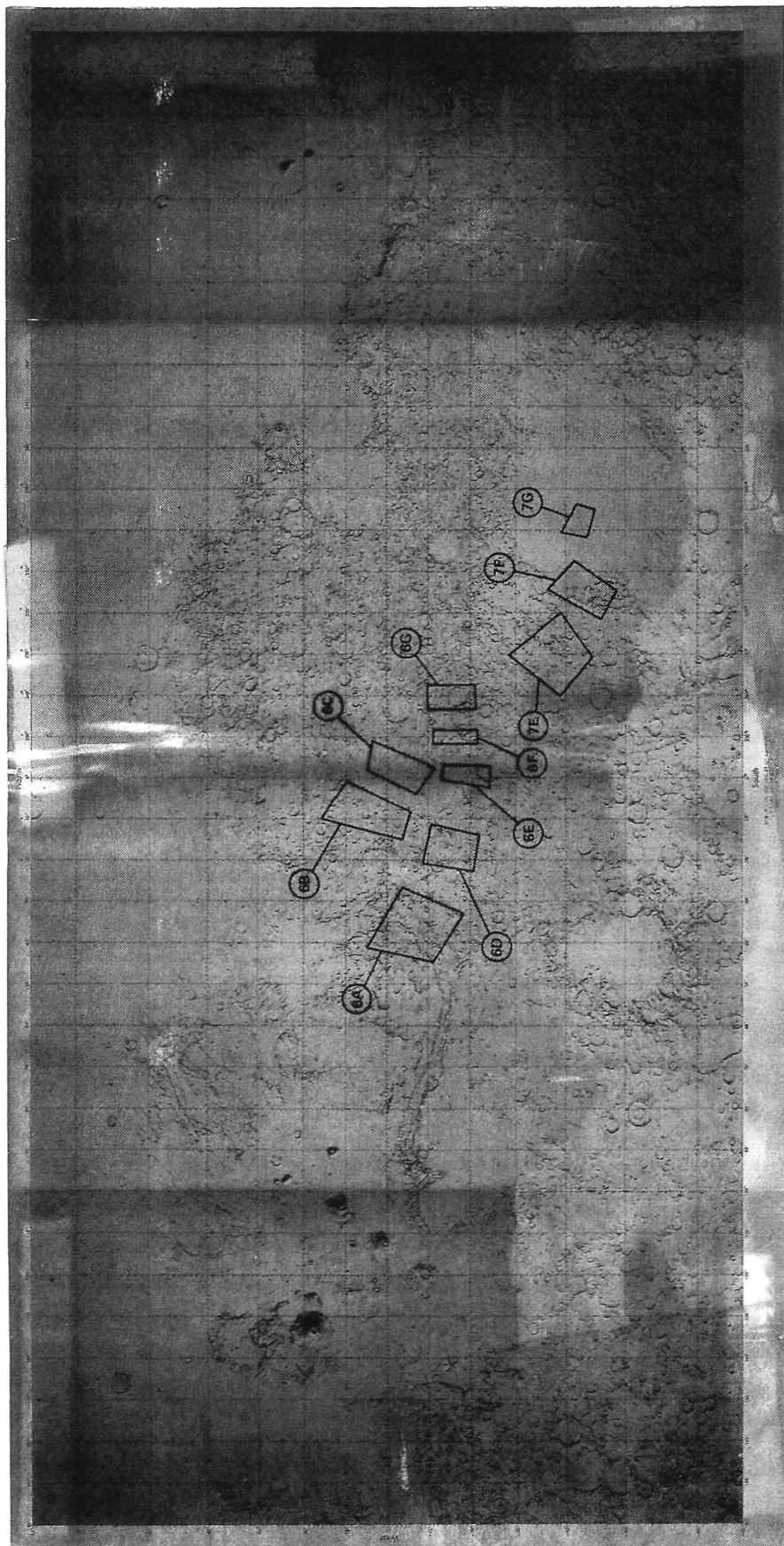


Figure 5 BI-BAND COVERAGE AREAS ON MARS

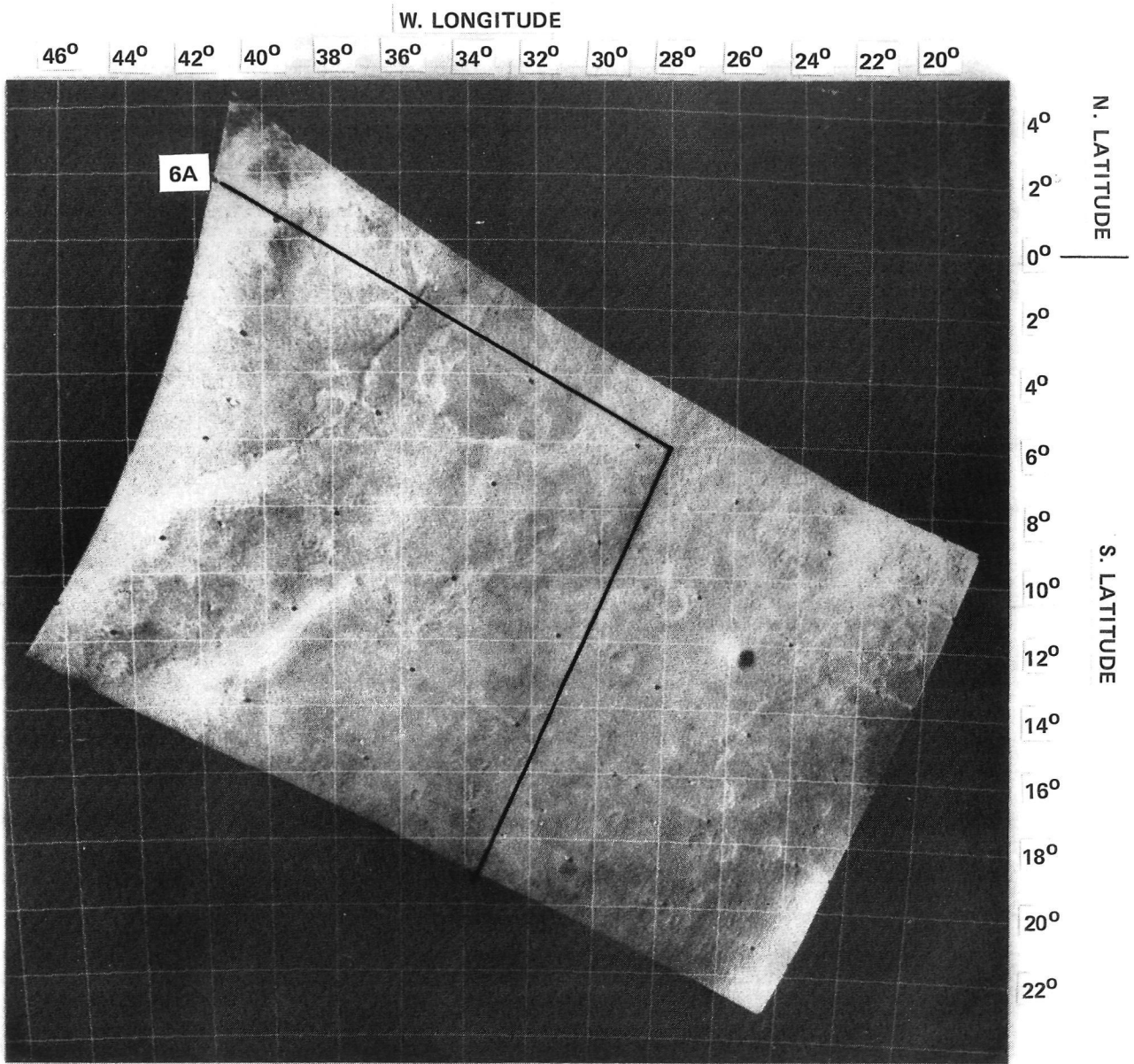


Figure 6 BI-BAND AREA 6A SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 6N7



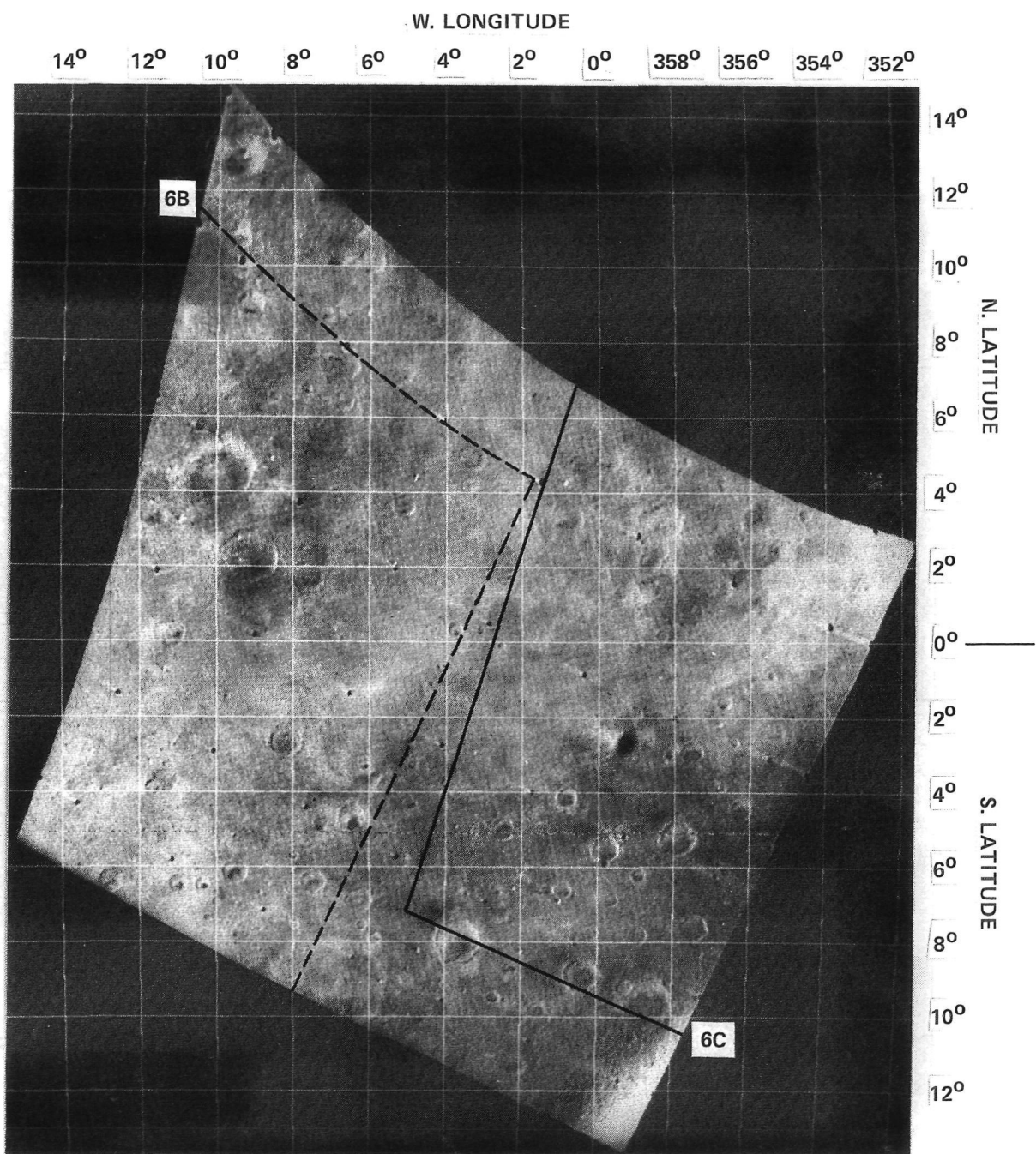


Figure 7 BI-BAND AREAS 6B AND 6C SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 6N11

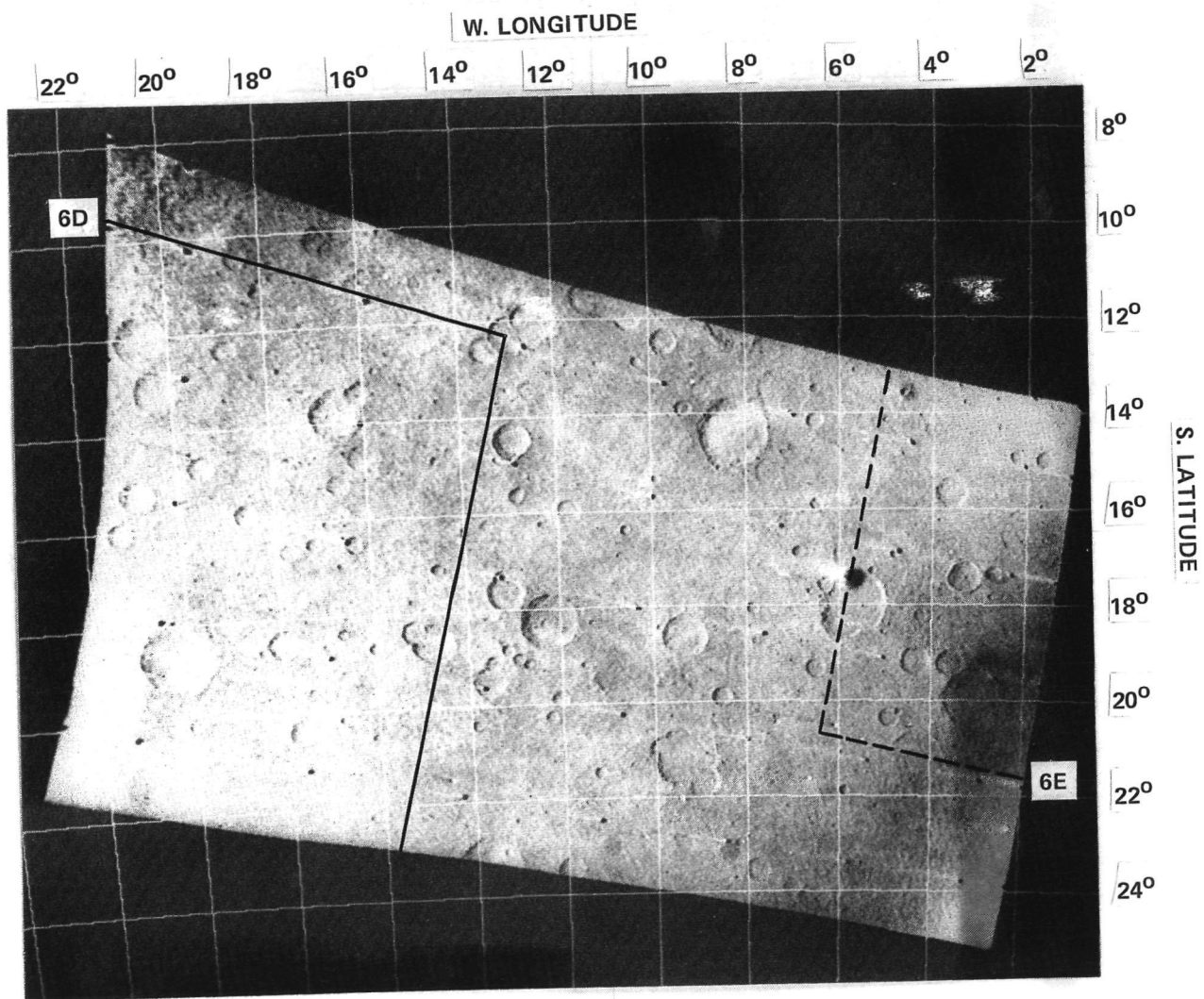


Figure 8 BI-BAND AREAS 6D AND 6E SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 6N17

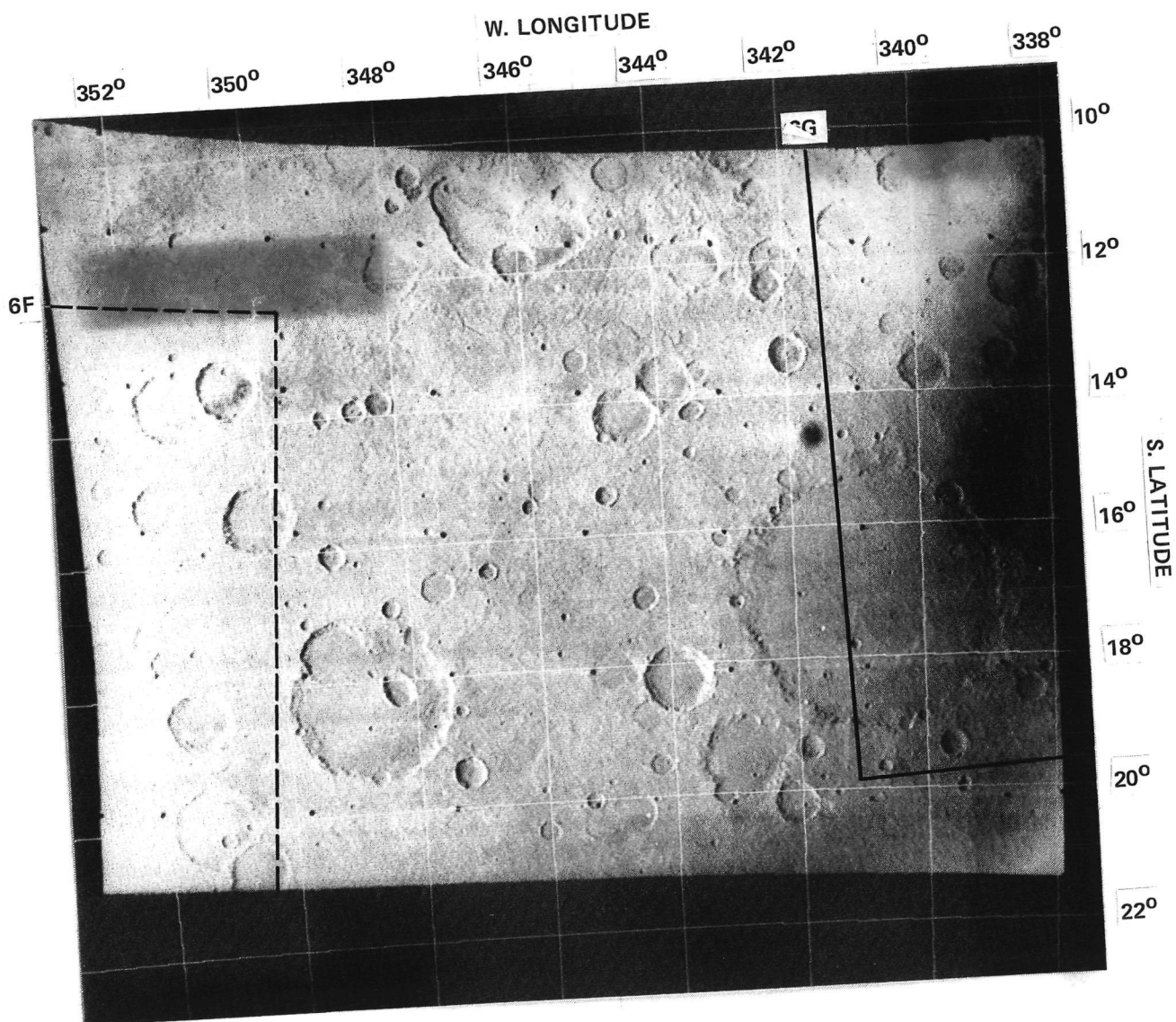


Figure 9 BI-BAND AREA 6F AND 6G SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 6N21

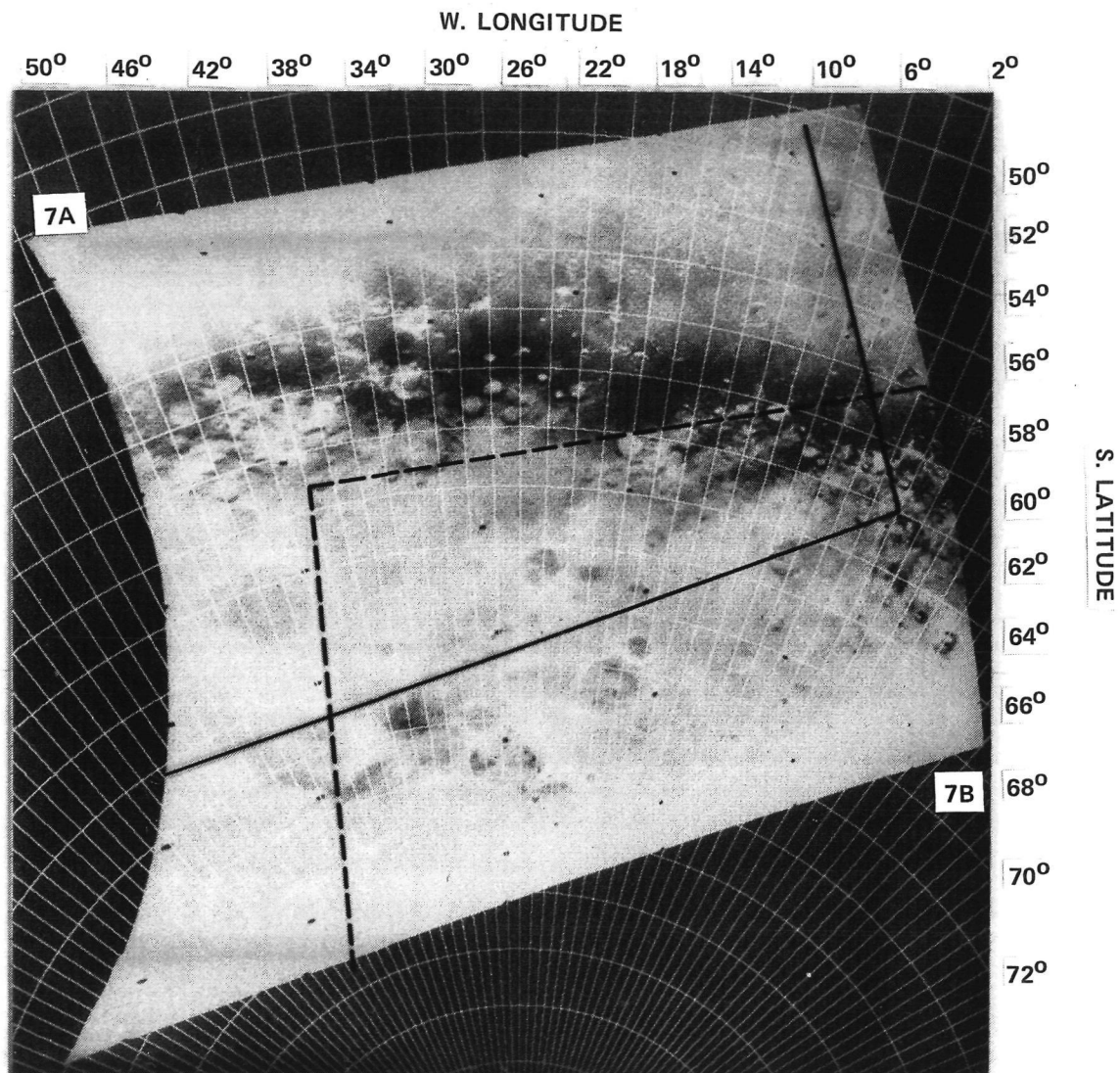


Figure 10 BI-BAND AREAS 7A AND 7B SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 7N13



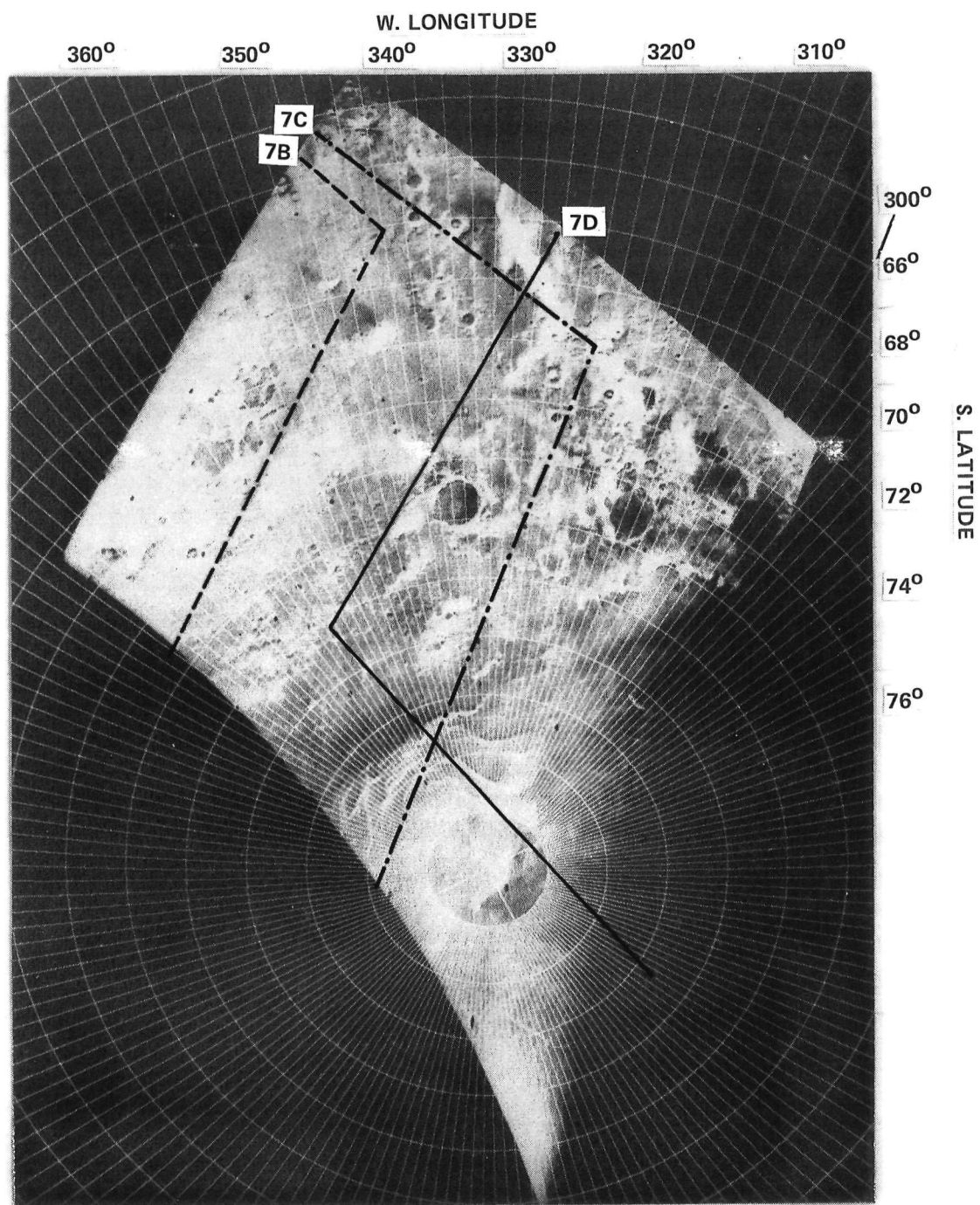


Figure 11 BI-BAND AREAS 7B, 7C AND 7D SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 7N17

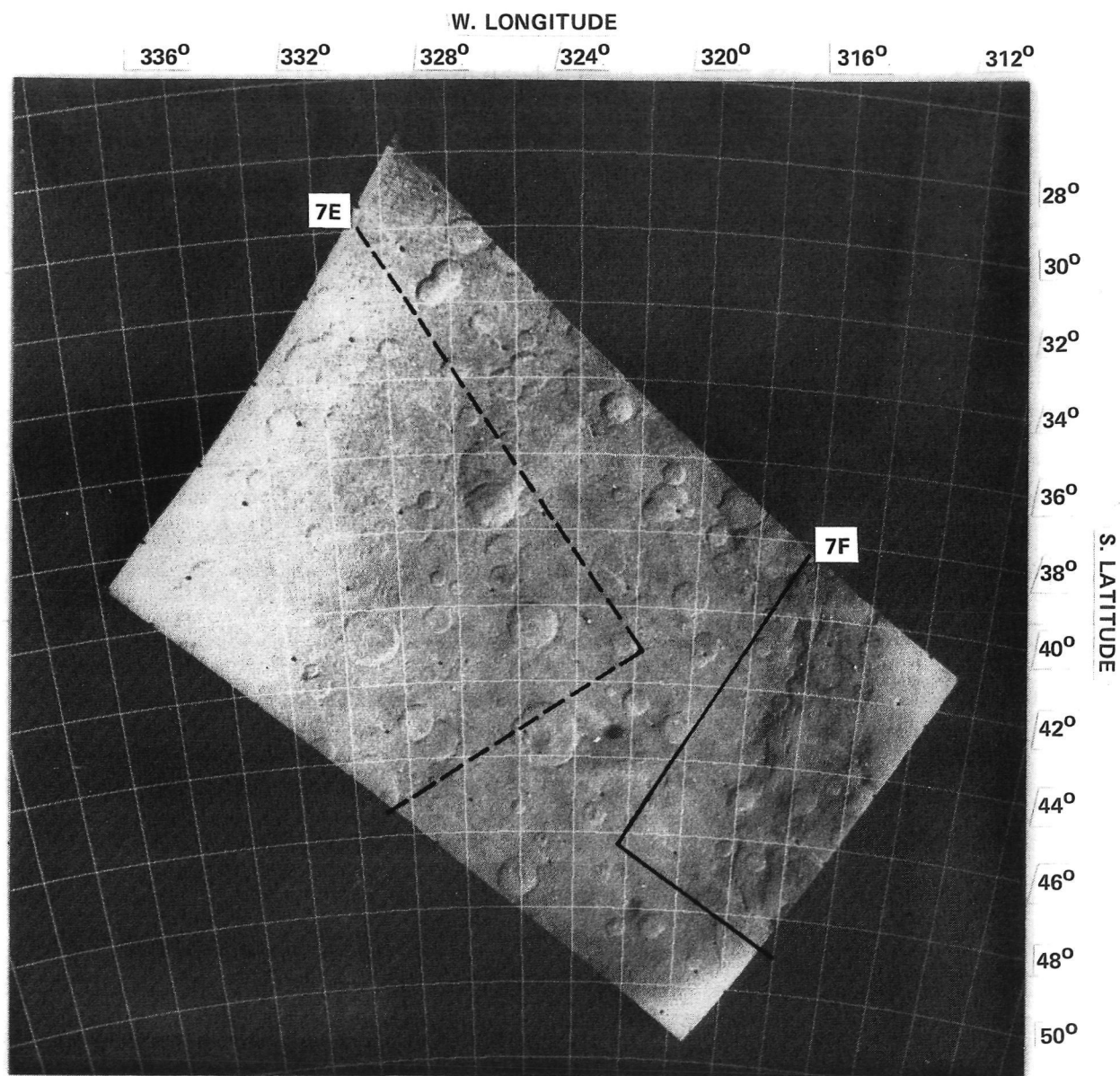


Figure 12 BI-BAND AREAS 7E AND 7F SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 7N25

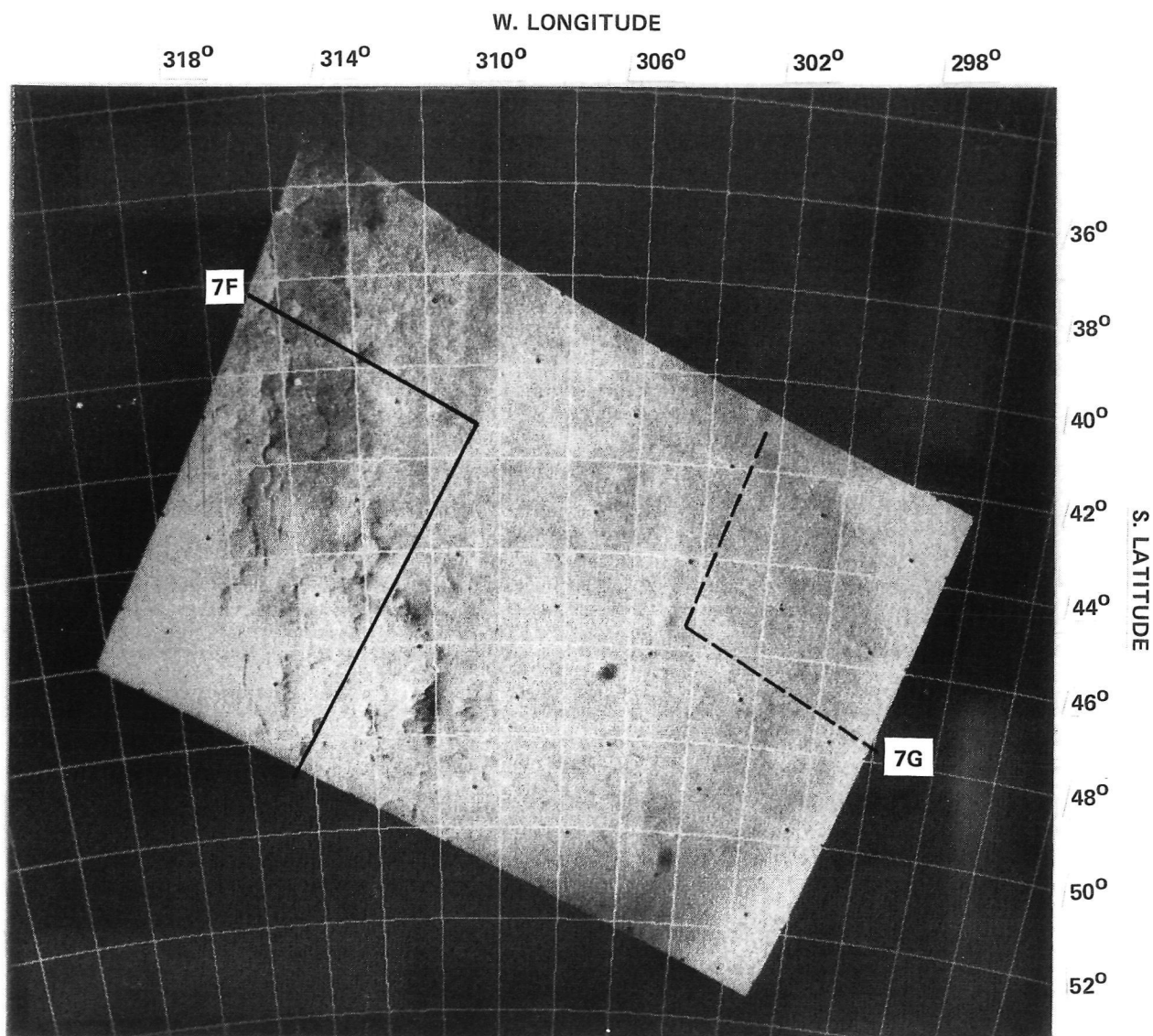


Figure 13 BI-BAND AREAS 7F AND 7G SHOWN ON ORTHOGRAPHIC PROJECTION OF FRAME 7N27

Jet Propulsion Laboratory. The data obtained was the LMICOR version which was produced by the Image Processing Laboratory at JPL by correcting the raw TV data for radiometric distortion introduced by the TV systems. The corrected data represents the surface bidirectional reflectivity multiplied by a factor of 6.0 in order to optimally utilize an 8-bit format.

The area that received the most detailed analysis was 6C. This area has light and dark markings on crater floors which were evaluated to determine whether this variation is caused by absorbed water or by changes in particle size. The method employed uses the reflectivity ratios in the manner discussed in Section 4.3. Section 5.2 contains additional data on the relationship between particle size and reflectivity that supports our analysis of these markings. The results of the analysis itself are contained in Section 5.4.

## 5.2 Reflectivity - Particle Size Effect

A size-reflectivity change rule was described in Section 4.3 that has been utilized in terrestrial soil surveys. This rule indicated that the reflectivity ratio of the dark-to-light soil element should decrease with increasing wavelength if a change in particle size was responsible for the change in reflectivity.

The proposal that limonite be a major constituent of the martian surface led Salisbury and Hunt (1968) to conduct a study of the spectral behavior of likely martian surface materials. They classified potential



materials by their change in reflectivity as a function of particle size. The class of transparent materials, which includes silicate minerals, show a increase in reflectivity with a decrease in particle size as assumed in the development of the size-reflectivity change rule referenced above. Opaque material, particularly metal sulfides, showing a decrease in reflectivity with decreasing particle size. A third class of materials which Salisbury and Hunt term "trans-opaque" exhibits a decrease in reflectivity in the opaque portion of the visible spectrum and an increase in reflectivity in the transparent portion of their visible spectrum. That is, the spectral reflectivity curve for one particle size will cross that for a different particle size leading to a reversal in reflectivity between the two particle sizes within the visible spectrum. This behavior is exhibited by several ferric oxides such as limonite, goethite and hematite that are likely candidates for existence on the martian surface. As Salisbury and Hunt point out, light and dark markings on the martian surface composed of high concentrations of these types of materials should exhibit a contrast reversal at some point in the visual spectral region. That is, the light markings should become dark and the dark markings become light as the radiometric or image data increase in wavelength. Since such a reversal has not been observed Salisbury and Hunt concluded that the hypothesis that the martian soil consists in a large part of limonite and that the reflectivity differences are due to particle size changes are incompatible. As an alternative they propose that the most likely soil is one composed of silicates lightly stained or coated with ferric oxides. In terms of the present evaluation of the Mariner imagery the dark-to-light reflectivity ratio could increase with

wavelength between the blue and green images and subsequently decrease with wavelength between the green and red images if a significant amount of the trans-opaque type materials composed the elements under evaluation and the change in particle size was the principal cause for the change in reflectivity. Our analysis concentrated on the local reflectivity changes occurring in bi-band coverage area 6C involving overlapping green and red imagery.

To further support our analyses using the reflectivity ratio, the ratios for typical martian materials were evaluated over the spectral responses of the Mariner VI and Mariner VII wide angle cameras presented earlier in Figure 4. Hunt and Salisbury (1970, 1971) have obtained a number of spectral reflectivity curves for various rocks and minerals as part of a laboratory study of spectroscopic remote sensing techniques. We selected typical candidates from the published data for silicate minerals (1970) and oxides and hydroxides (1971). The spectral reflectivity for each of the minerals selected is shown in Figure 14 (although measurements were made at IR wavelengths this data was not of interest here). The reflectivity data were obtained near normal incidence and emission angles at a phase angle of  $15^\circ$  and referenced to a Lambertian surface with a diffuse reflectivity near 1.0. Generally the data were obtained for four different particle size regions allowing six different ratios of the reflectivity of larger-to-smaller particle sizes to be obtained. For materials whose reflectivity increases with decreasing particle size this ratio is equivalent to that of the darker-to-lighter soil elements. The six different ratios obtained are identified in Table 4 using the letter symbol A through F.

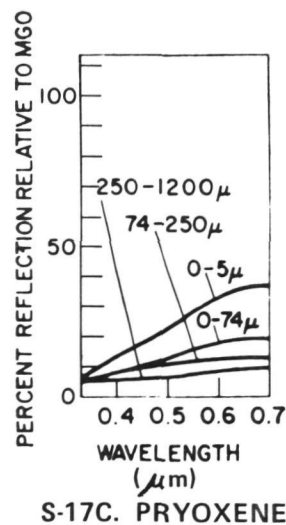
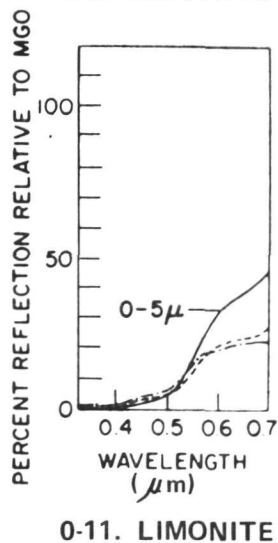
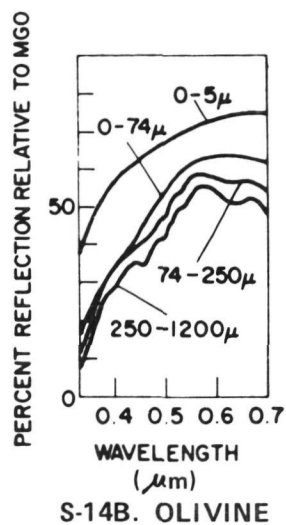
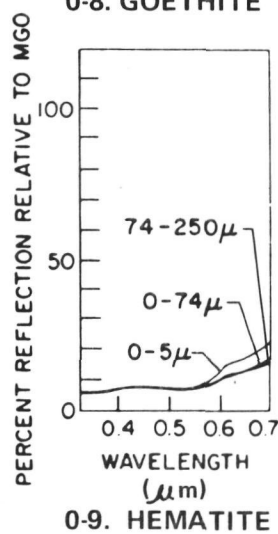
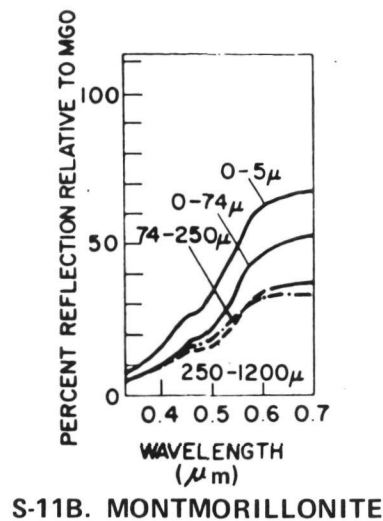
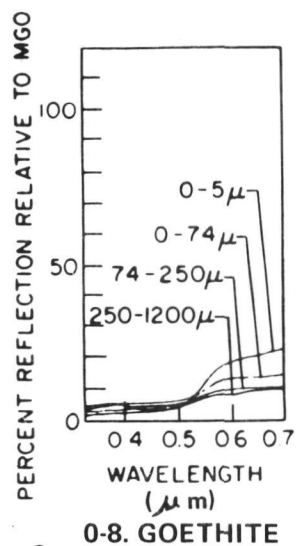


Figure 14 SPECTRAL REFLECTIVITY FOR TYPICAL MARTIAN MINERALS AS A FUNCTION OF PARTICLE SIZE

Table 4

## Large-to-Small Particle Index

Symbol	Particle Size Ratio
A	250-1200 $\mu\text{m}$ /0-5 $\mu\text{m}$
B	250-1200 $\mu\text{m}$ /0-74 $\mu\text{m}$
C	250-1200 $\mu\text{m}$ /74-250 $\mu\text{m}$
D	74-250 $\mu\text{m}$ /0-5 $\mu\text{m}$
E	74-250 $\mu\text{m}$ /0-74 $\mu\text{m}$
F	0-74 $\mu\text{m}$ /0-5 $\mu\text{m}$

From the silicate minerals olivine and pyroxene were selected based upon the analysis of Binder and Jones (1972). Montmorillonite was arbitrarily added as the third selection from the silicate minerals. Limonite was selected from the ferric oxides since it has been proposed as a major constituent by Pollack and Sagan (1967) and as a stain on silicates by Binder and Jones (1972) and Salisbury and Hunt (1968). Pollack and Sagan proposed goethite as a major constituent of the martian surface and consequently this was also selected by the oxide group. Finally, Hematite was added as a third ferric oxide--a logical addition to those already selected. Table 5 presents the average reflectivity of these minerals within the Mariner VI and VII spectral bands. As the table shows, the reflectivity increases with decreasing particle size in all cases except for the smallest particle size group of goethite in the green band (and presumably the blue band) and for all of the particle size groups measured in the case of limonite for both the blue and green spectral bands. Consequently, the ratio of the larger-to-smaller particle size groups obtained from this data are equivalent to the ratio of the darker-to-lighter surface elements except for the cases cited.

Table 5

Average Reflectivity of Potential  
Martian Materials Within the Mariner Spectral Bands

Mineral	Size Range	Mariner VI			Mariner VII		
		Blue	Green	Red	Blue	Green	Red
0-8 Goethite	250-1200 $\mu$ m	.048	.064	.083	.049	.064	.085
	74- 250 $\mu$ m	.039	.065	.089	.037	.066	.092
	0- 74 $\mu$ m	.060	.080	.116	.059	.081	.121
	0- 5 $\mu$ m	?	.078	.152	?	.079	.162
0-9 Hematite	74- 250 $\mu$ m	.072	.068	.082	.072	.068	.087
	0- 74 $\mu$ m	.080	.075	.089	.081	.075	.093
	0- 5 $\mu$ m	.081	.080	.103	.081	.080	.110
0-11 Limonite	250-1200 $\mu$ m	.054	.104	.174	.053	.106	.182
	0- 74 $\mu$ m	.040	.088	.164	.040	.089	.177
	0- 5 $\mu$ m	.030	.097	.228	.029	.099	.250
S-11B Mont- morillonite	250-1200 $\mu$ m	.144	.201	.292	.143	.203	.306
	74- 250 $\mu$ m	.166	.227	.290	.162	.229	.298
	0- 74 $\mu$ m	.186	.276	.404	.182	.279	.422
	0- 5 $\mu$ m	.270	.407	.547	.265	.411	.568
S-14B Olivine	250-1200 $\mu$ m	.373	.478	.536	.365	.481	.539
	74- 250 $\mu$ m	.427	.530	.577	.418	.533	.578
	0- 74 $\mu$ m	.473	.573	.617	.459	.576	.620
	0- 5 $\mu$ m	.637	.681	.713	.631	.682	.718
S-17C Pyroxene	250-1200 $\mu$ m	.058	.066	.077	.057	.066	.079
	74- 250 $\mu$ m	.098	.113	.120	.096	.114	.121
	0- 74 $\mu$ m	.108	.137	.162	.105	.138	.165
	0- 5 $\mu$ m	.193	.254	.305	.188	.256	.312

Figure 15 shows the reflectivity ratios determined for the ferric oxide samples and the Mariner VI spectral response functions. The ratios were plotted at the effective focal length for each of the spectral filters as specified by Danielson (1971). Figure 16 is the corresponding plot for the silicate samples. The results obtained in the case of Mariner VII are similar to the reflectivity ratios shown here. We would expect that the ratios would approach unity as the particle size differences between the two groups being ratioed decreases. In addition, the rate of change of the ratio with respect to wavelength should decrease as the particle size differences decrease, i.e., the slope of the curve becomes less. If we compare the progression of curves for cases A to B to C we would expect that these phenomena would be apparent in Figures 15 and 16. Except for the cases where a reversal in the change of reflectivity with particle size occurs, this phenomena is exhibited for all of the minerals evaluated. A similar trend is exhibited in going from the curves with index D to the curves of index E. We can infer from these results that information about the relative difference in particle sizes can be obtained by examining the level and slope of the reflectivity ratio versus wavelength curve. The difference between the ratio from the green to the red band is presented in Table 6. In this table a positive number indicates a decrease in the ratio as the wavelength is increased as expected from the size-reflectivity change rule stated previously. Note that most of the changes are positive (i.e., show a decrease with increasing wavelength) except for olivine which shows an increase in the ratio for all cases. If we eliminate this sample from consideration the average value of the change which is also presented in the table is greatest for the A ratio which has

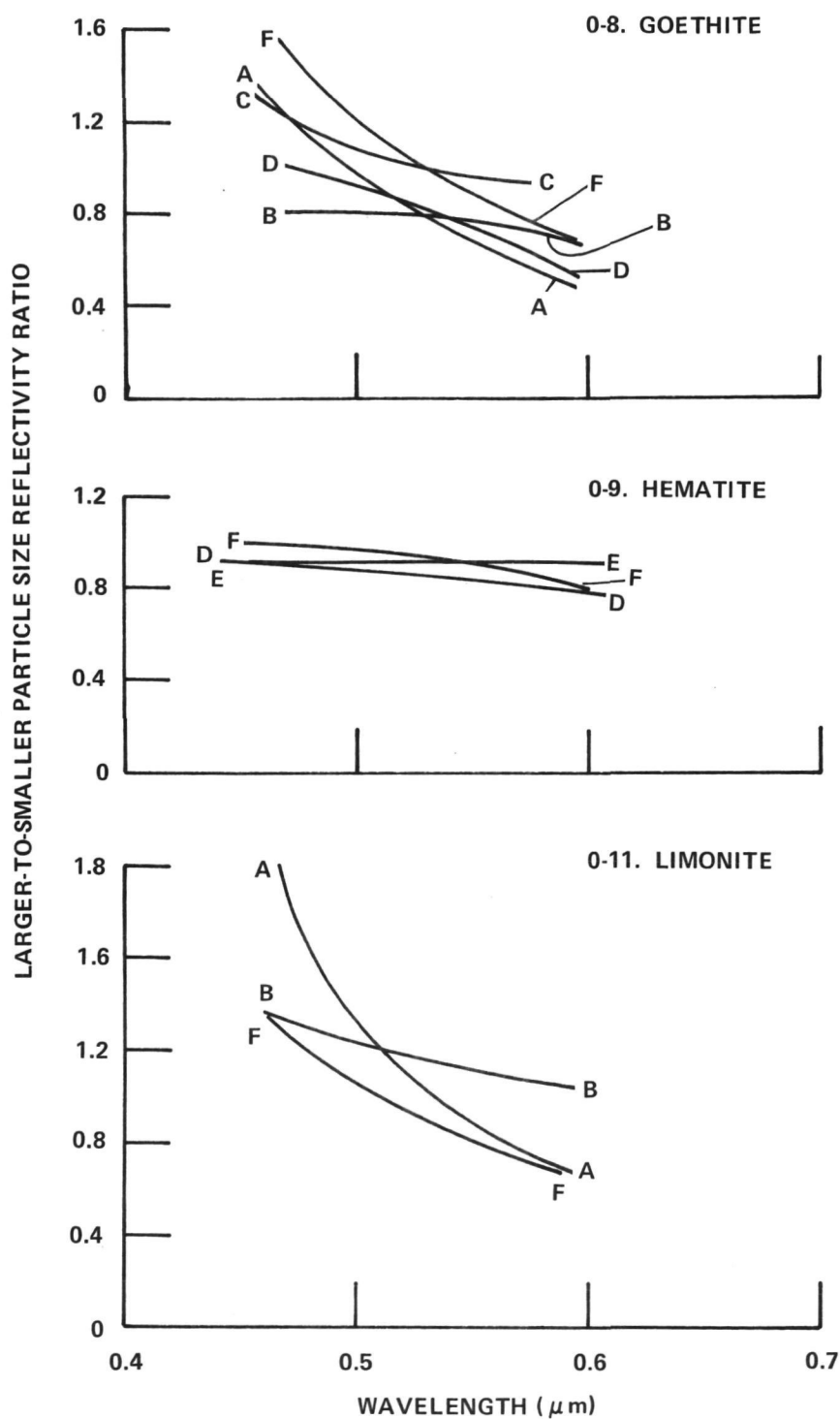


Figure 15 REFLECTIVITY RATIOS FOR SELECTED OXIDES AND VARIOUS PARTICLE SIZE GROUPS

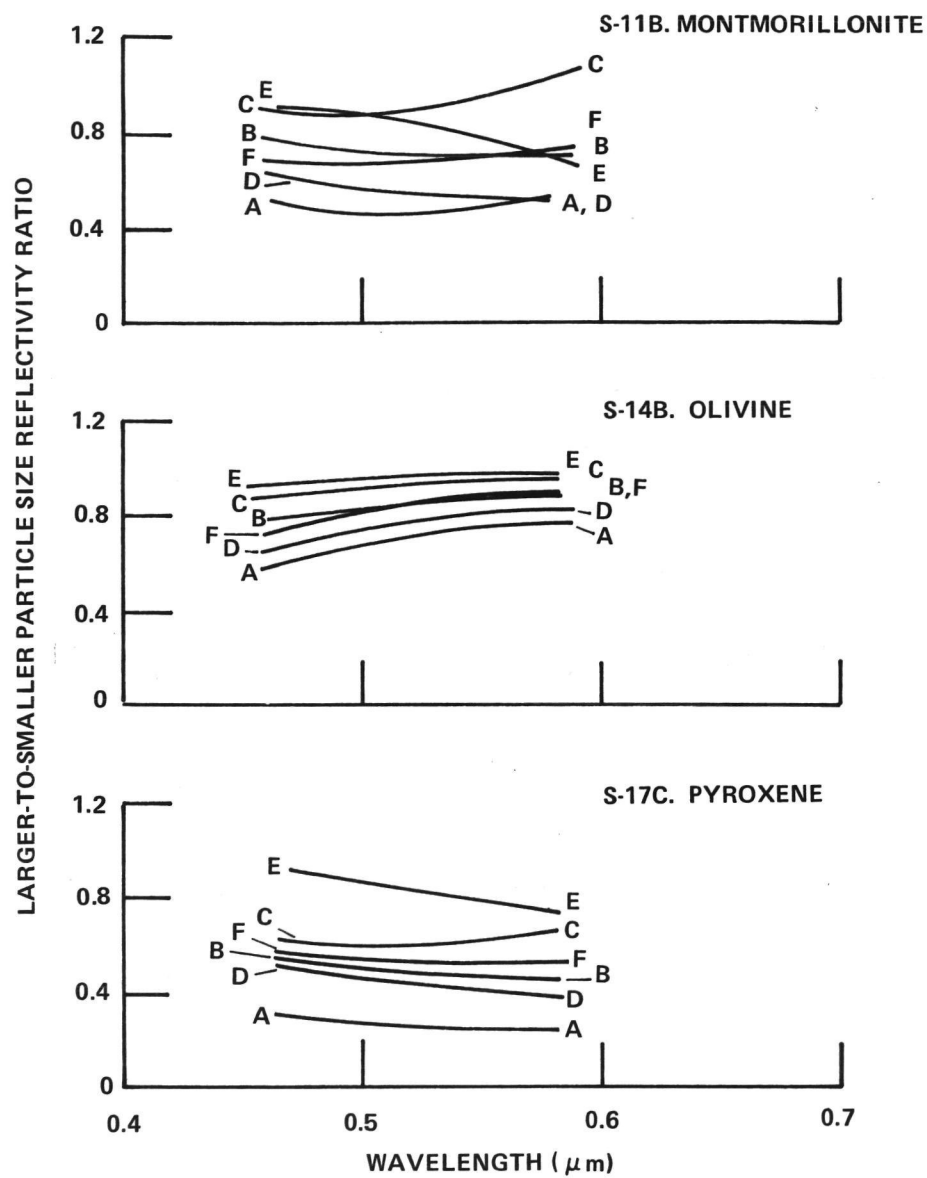


Figure 16 REFLECTIVITY RATIOS FOR SELECTED SILICATES AND VARIOUS PARTICLE SIZE GROUPS



Table 6

Green-to-Red Decrease in the Reflectivity Ratio

Index	0-8 Goethite	0-9 Hematite	0-11 Limonite	S-11B Montmorillonite	S-14B Olivine	S-17C Pyroxene	Average (w/o Olivine)
A	.27	-	.31	-.04	-.03	.01	.14
B	.08	-	.12	.01	-.04	.00	.05
C	.05	-	-	-.11	-.05	-.06	-.04
D	.24	.05	-	.03	-.02	.05	.09
E	.04	-.01	-	.10	-.03	.08	.05
F	.30	.08	.19	-.06	-.03	.01	.10

the largest particle size differences and least (negative) for the C ratio where the particle size differences are not as great. Consequently, we conclude that the slope of the reflectivity ratio versus wavelength curve indeed does contain information about the relative particle size of the materials producing the light and dark markings.

### 5.3 Evaluation of the Light and Dark Markings Associated with Craters

Bi-band coverage area 6C contained several craters with light and dark markings located interior to the crater on their floor. Several of these craters were chosen for subsequent analysis of the reflectivity ratio of these markings from frames 6N11 and 6N13. The craters selected are shown in Figure 17. The light and dark markings were located by sample and line number and subsequently retrieved from the LMICOR images contained on digital tape supplied by JPL. Software programs were developed to retrieve this data over a rectangular area specified by the initial and final sample and line number. A description of this software is presented in the appendix to this report. Table 7 contains the average reflectivity and subsequent ratio for each of the selected dark and light markings in the red and green spectral bands.

These data are not the only source of reflectivity ratios of dark-to-light markings on the martina surface. Dollfus measured the dark-to-light area contrasts in several wavelength bands. These data were employed by Pollack and Sagan (1967) and were extracted for comparison to our results. These results are presented in Table 8. Other studies utilizing reflectivity

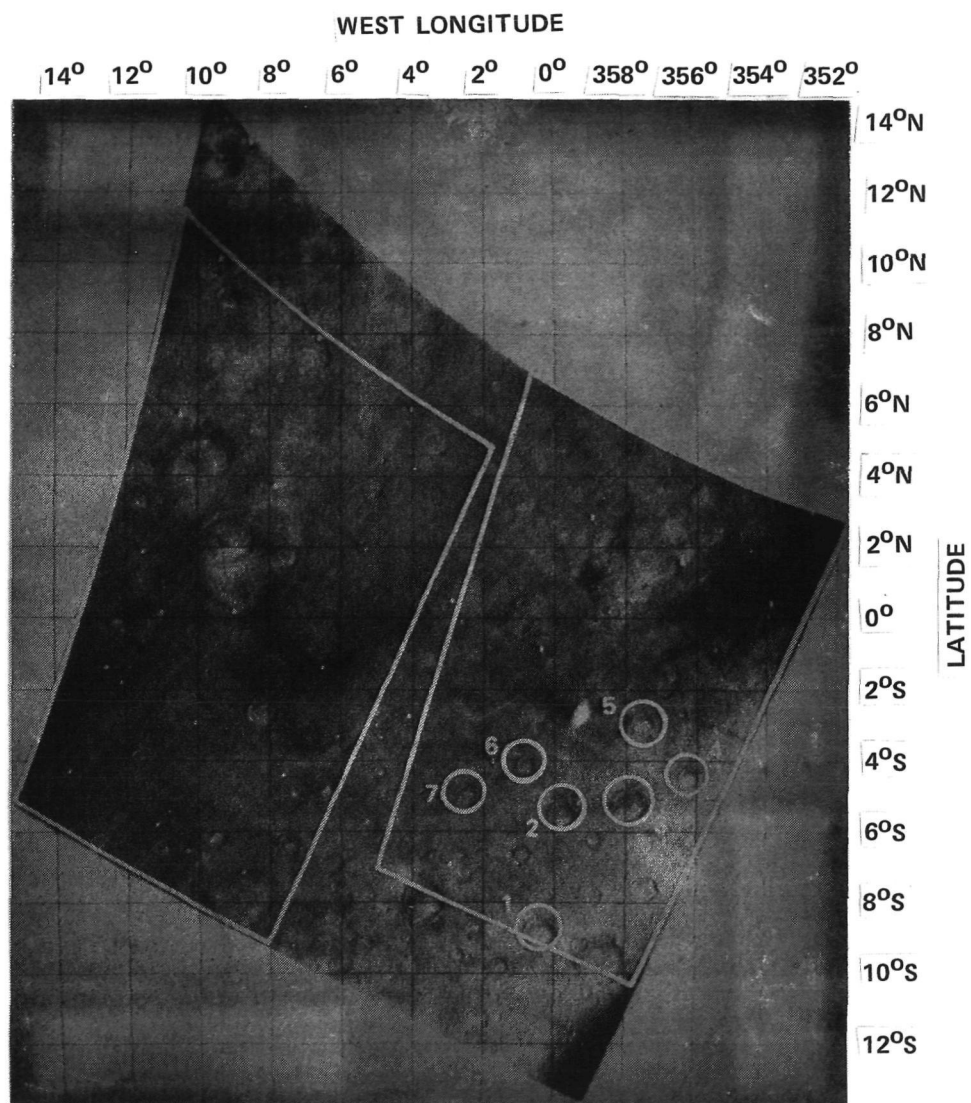


Figure 17 CRATERS ANALYZED IN BI-BAND COVERAGE AREA 6C

Table 7

## Reflectivity and Reflectivity Ratio of Crater Markings

Target	6N11-Green (.53 $\mu$ m)			6N13-Red (.58 $\mu$ m)		
	Dark	Light	Ratio	Dark	Light	Ratio
1 - Crater	.0588	.0602	.977	.0742	.0879	.844
2 - Crater	.0583	.0633	.921	.0708	.0843	.840
3 - Crater	.0592	.0640	.925	.0658	.0816	.806
4 - Crater	.0604	.0632	.956	.0668	.0754	.886
5 - Crater	.0598	.0641	.933	.0665	.0767	.867
6 - Crater	.0603	.0651	.926	.0717	.0806	.890
7 - Crater	.0599	.0638	.939	.0734	.0764	.961
Average	.0595	.0634	.940	.0699	.0804	.871
$\sigma$	.0009	.0015	.020	.0035	.0046	.049

$\Delta$  = Green-to-Red Change = .07

Table 8

## Normal Reflectivity and Dark-to-Light Reflectivity Ratios Derived from Dollfus Contrast Data

$\lambda$ ( $\mu$ m)	Dark Areas	Bright Areas	Ratio
.45	.065	.071	.92
.50	.098	.120	.82
.55	.120	.164	.73
.60	.139	.210	.66
.65	.150	.250	.60

$\Delta$  = Effective Green-to-Red Change = .09

ratios were conducted by McCord (1969) and by Cutts (1971). McCord studied selected dark and bright regions primarily the Arabia-Syrtus Major pair at 21 narrow spectral passbands in the visible spectrum. From his results, McCord concluded that the bright regions are much redder than the dark regions, that is the ratio of the reflectivity of the dark-to-light regions in the red region is greater than the blue region. According to the size-reflectivity change rule, this would indicate that particle size effects are probably responsible for the differences between the bright and dark martian regions. Cutts subsequently employed the McCord data to augment his evaluation which was obtained using the late far encounter images obtained during the Mariner VII mission. Cutts evaluated nine different regional areas ranging from high to low reflectivity. The reflectivity ratio for dark-to-light areas for selected data presented by Cutts is contained in Table 9. The first three ratios were selected to represent that of the darkest-to-lightest features. The second two ratios were selected because they represent moderately dark-to-light features. Note that the value of the reflectivity ratio is closer to one for the second group than the first group and that the change or average slope of the reflectivity ratio curve is less for the second group compared to the first group. The third group contains two cases, the first being the ratio between two relatively dark features and the second being the ratio between two relatively light features. Note in this case that the reflectivity ratios are much closer to unity and that the average slope is less than either of the two preceding groups.

Table 9

Dark-to-Light Reflectivity Ratios Derived from  
Selected Cutts/McCord Data Within Mariner Spectral Bands

	Description	Blue (.47 $\mu$ m)	Green (.53 $\mu$ m)	Red (.58 $\mu$ m)	(Green-to-Red)
I	Syrtis Major/Arabia*	.87	.77	.65	.12
	Margaritifer Sinus/Moab	.83	.73	.59	.14
	Meridiani Sinus/Moab	.84	.75	.65	.10
II	Deucalionis Regio/Moab	.97	.95	.93	.02
	Thymiamata/Moab	.92	.87	.80	.07
III	Meridiani Sinus/Sabaeus Sinus	.95	.94	.90	.04
	Edom/Moab	1.01	1.00	.97	.03

\*After Cutts

All three data sources were plotted for comparison to the Mariner VI crater markings in Figures 18 and 19. Figure 18 shows that the crater markings show a change in ratio which indicates that particle size differences are responsible for the change in reflectivity (the ratio decreases with increasing wavelength). Furthermore, comparison of the Mariner results to that of Cutts and McCord indicates that the Mariner ratios represent more moderate changes in particle size than that due to the extreme reflectivity changes. The change in reflectivity ratio from green-to-red of 0.07 is compatible with those in Group II in Table 9 and is consistent with the changes obtained in Section 4.2 for particle size groups B and E. A similar result is exhibited in Figure 19 when compared with the Dollfus ratios. The Dollfus results appear to represent a larger change in reflectivity than that due to the crater markings. The

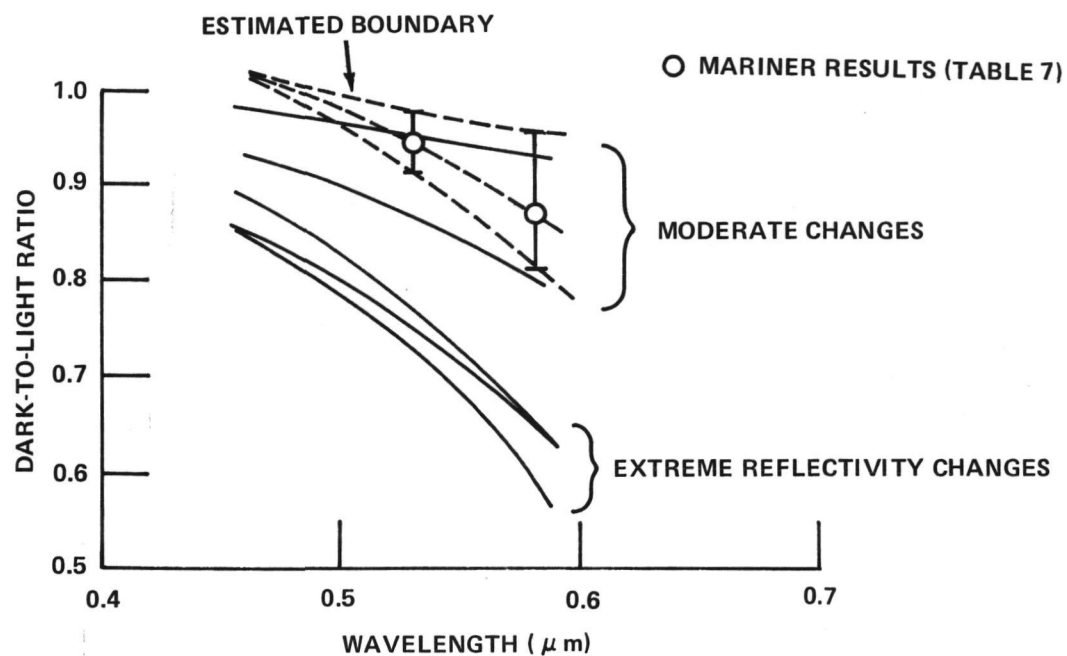


Figure 18 COMPARISON OF MARINER REFLECTIVITY RATIOS TO THOSE OF CUTTS AND McCORD

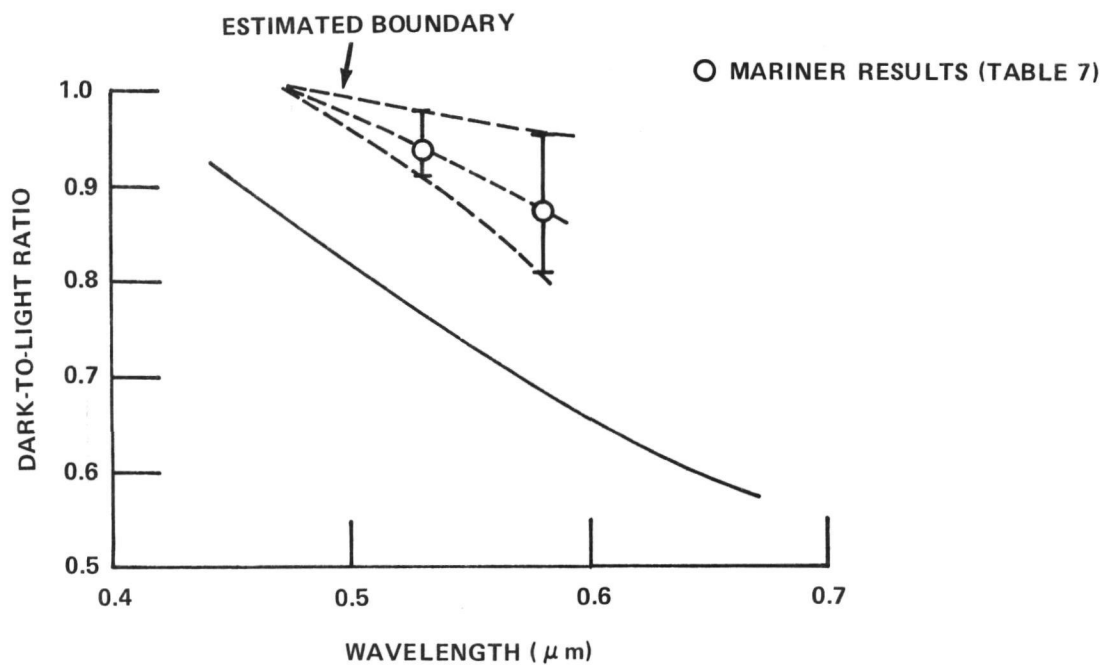


Figure 19 COMPARISON OF MARINER RATIO TO THOSE OF DOLLFUS



Group I data of Cutts and McCord (Table 9) and that of Dollfus (Table 8) are consistent with size Groups A and D in Section 4.2. Note that the size ranges represented by Groups B and E are compatible with the postulated mean particle sizes for the light and dark areas offered by Pollack and Sagan (1967).

If we compare the reflectivity presented in Table 5 to an extrapolation of the Minnaert parameter  $N_g$  obtained by Thorpe (Table 2) we find that the normal reflectivity of olivine appears to be a factor of 2-3 greater than the average normal albedo. Consequently, limonite stained olivine is not considered to be a likely candidate as a major constituent on this local area.

Limonite stained pyroxene appears to be the best candidate based upon the data obtained during the course of this effort. Of course this conclusion must be somewhat tenuous since a number of mineral compositions could yield the results obtained.

#### 5.4 Mariner IX Imagery and Other Related Data

We had hoped to continue the present effort using Mariner IX image data obtained through multiple spectral filters. It was anticipated that a large number of multi-spectral coverages would be available. Unfortunately, the filter wheel mechanism became inoperative during Revolution 118 and only a very limited amount of multi-spectral data was obtained. For the subsequent portion of the Mariner mission image data was obtained using the orange polarizing filter contained in the fifth filter wheel position. During the initial part of the mission the martian surface was obscured by a dust storm and consequently no useful image data of surface features were obtained in

this part of the mission when the filter wheel was operative. Mariner IX did confirm that dynamic aeolian processes play a significant role in the reflectivity variations on the martian surface and that a significant number of these variations are caused by changes in particle size.

An analysis of the light and dark markings present in the Mariner IX imagery obtained subsequent to the subsiding of the dust storm was made by Arvidson (1974). As a result of his analysis he found that "dark spotted craters in regions with bright streaks usually have upwind bright patches, suggesting that these features formed by dumping of bright dust over crater rims ... ". This characteristic pattern minus the bright streaks downwind of the craters are identical to those that we have analyzed. Furthermore, they have a North to South orientation consistent with the direction of the wind pattern observed subsequent to the dust storm through analysis of the wind-blown streaks. In addition, a sequence of wind tunnel experiments have been carried out by Greeley and his associates (1974). The results of these experiments have generated light and dark markings similar to those observed during our effort and have further demonstrated that some of the dark markings occur from wind erosion and that some of the light markings are depositional in nature. The reader is also referred to a comparison of Mariner VI and VII imagery and Mariner IX imagery which includes the bi-band coverage area we analyzed for a qualitative assessment of effect of the dust storm on local reflectivity variations (Ververka, 1974).

Though it was intended to demonstrate the use of reflectivity ratios as an aid to geologic mapping, the lack of availability of Mariner IX imagery greatly degraded this objective. The Mariner VI and VII imagery within the bi-band coverage areas occurred over a limited number of geologic units and the detailed geologic maps for the Quadrangles containing these areas had not as yet been produced. It was also intended to obtain information about the martian atmosphere through radiometric measurements of the shadows associated with geological features, especially bowl-shaped craters. Although a number of shadows were identified in the narrow field of view imagery from Mariner VI, in particular Frame 6N22, the lack of available shadows is a direct result of the large diameter/depth ratios observed for Mars compared to Mercury and the Moon (Cintala, 1976). The diameter to depth ratio for craters near 10 km in diameter is approximately 8 times larger while for craters near 100 km in diameter it is approximately 3 times larger. The lack of Mariner IX bi-band image data further inhibited the evaluation of atmospheric phenomena.

## SECTION 6

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## APPENDIX

### ANALYSIS SOFTWARE

This appendix describes three items of analysis software which were prepared during the course of this study. The principal software package consisted of the main program written in Fortran language and an associated subroutine, EPIC, written in Assembly language. Listings of the main program and subroutines programs are included in Figures A-1 and A-2 respectively. The purpose of this program is to retrieve a rectangular array of reflectance data from digital tapes of Mariner imagery provided by JPL and compute the mean, standard deviation, minimum and maximum of the reflectivities in the array. The data required is retrieved from the digital tapes using the Subroutine FPIC and subsequently analyzed by the main program whose listing is included in Figure A-1. In addition to computing the basic statistics, this main program also has the option of computing a histogram of the reflectivity values in the specified array.

Subroutine FPIC retrieves each element in a rectangular array located within the digital TV picture data for subsequent processing by the main program. The rectangular area of interest is defined by the following calling sequence:

CALL FPIC (NLS, NSS, NSE, NOL, ARRAY, NLA, NSA, RFAC) where:

NLS: is the line number within the picture where data retrieval begins

NSS: is the sample number within the line where data retrieval begins

```

C      THIS PROGRAM CALCULATES THE STATISTICS OF A RECTANGULAR IMAGE ARRAY
C      SUPPLIED BY SUBROUTINE FPIC.
C
ISN 0002      DIMENSION R(25,25),IHIST(15)
ISN 0003      REAL*8 TARGID
ISN 0004      DATA IOPT/0/
ISN 0005      1 FORMAT(A8,50X,4I4,F4.0,I2)
ISN 0006      9 FORMAT(1H1)
ISN 0007      10 FORMAT(6X 'TARGET IDENTIFICATION= ',A8,'      INITIAL LINE NO.=' ,I4,
      *' INITIAL SAMPLE NO.=' ,I4,'      FINAL SAMPLE NO.=' ,I4,'      NO. OF LINE
      *S= ',I4 /// 13X 'PERCENT REFLECTANCE: MINIMUM= ',F6.3,'      MAXIMUM= ',
      *F6.3,'      AVERAGE= ',F6.3,'      STANDARD DEVIATION= ',F8.5 ///
      *13X 'HISTOGRAM: CLASS INTERVAL= ',F6.3, // 16X 'CLASS NO.      ',
      *15(I3,1X) / 16X 'NO. OF PIXELS ',15(I3,1X) / 16X 'TOTAL NO. OF P
      *IXELS= ',I5 //)
ISN 0008      11 FORMAT(6X 'TARGET IDENTIFICATION= ',A8,'      INITIAL LINE NO.=' ,I4,
      *' INITIAL SAMPLE NO.=' ,I4,'      FINAL SAMPLE NO.=' ,I4,'      NO. OF LINE
      *S= ',I4 /// 13X 'PERCENT REFLECTANCE: MINIMUM= ',F6.3,'      MAXIMUM= ',
      *F6.3,'      AVERAGE= ',F6.3,'      STANDARD DEVIATION= ',F8.5 // 13X 'TOTAL
      *NO. OF PIXELS= ',I5 //)
ISN 0009      ICASE=-1
ISN 0010      NLA=25
ISN 0011      NSA=25
ISN 0012      50 READ(5,1,END=100) TARGID,NLS,NSS,NSE,NOL,RFAC,IOPT
ISN 0013      ICASE=ICASE+1
C
C      OBTAIN REFLECTANCE DATA ARRAY 'R'
C
ISN 0014      CALL FPIC(NLS,NSS,NSE,NOL,R,NLA,NSA,RFAC)
C
C      COMPUTE THE AVERAGE, STANDARD DEVIATION, MINIMUM AND MAXIMUM OF
C      THE 'R' ARRAY
C
ISN 0015      SUM=0.0
ISN 0016      SSQ=0.0
ISN 0017      RMIN=100.
ISN 0018      RMAX=0.0
ISN 0019      NOS=NSE-NSS+1
ISN 0020      NTOT=NOS*NOL
ISN 0021      DO 60 I=1,NOL
ISN 0022      DO 60 J=1,NOS
ISN 0023      RT=R(J,I)
ISN 0024      IF(RT.GE.RMIN) GO TO 55
ISN 0026      RMIN=RT
ISN 0027      55 IF(RT.LE.RMAX) GO TO 59
ISN 0029      RMAX=RT
ISN 0030      59 SUM=SUM+RT
ISN 0031      60 SSQ=SSQ+RT**2
ISN 0032      RAVE=SUM/FLOAT(NTOT)
ISN 0033      RSTD=SQRT((SSQ-FLOAT(NTOT)*RAVE**2)/FLOAT(NTOT-1))
ISN 0034      IF(IOPT.EQ.0)GO TO 80
C
C      COMPUTE HISTOGRAM
C
ISN 0036      NI=RFAC*(RMAX-RMIN)/15.
ISN 0037      CINT=FLOAT(NI+1)/RFAC
ISN 0038      DO 65 I=1,15
ISN 0039      65 IHIST(I)=0
ISN 0040      DO 70 I=1,NOL
ISN 0041      DO 70 J=1,NOS
ISN 0042      N=(R(J,I)-RMIN)/CINT +1.0
ISN 0043      70 IHIST(N)=IHIST(N)+1
ISN 0044      IF(MOD(ICASE,4).EQ.0)WRITE(6,9)
ISN 0046      WRITE(6,10) TARGID,NLS,NSS,NSE,NOL,RMIN,RMAX,RAVE,RSTD,CINT,
      *(I,I=1,15),(IHIST(I),I=1,15),NTOT
ISN 0047      GO TO 50
ISN 0048      80 IF(MOD(ICASE,4).EQ.0)WRITE(6,9)
ISN 0050      WRITE(6,11) TARGID,NLS,NSS,NSE,NOL,RMIN,RMAX,RAVE,RSTD,NTOT
ISN 0051      GO TO 50
ISN 0052      100 STOP
ISN 0053      END

```

Figure A-1 IMAGE STATISTICS MAIN PROGRAM LISTING



SYMBOL	INTERNAL STATEMENT NUMBERS											
I	0021	0023	0038	0039	0040	0042	0046	0046	0046	0046	0046	0046
J	0022	0023	0041	0042								
N	0042	0043	0043									
R	0002	0014	0023	0042								
NI	0036	0037										
RT	0023	0024	0026	0027	0029	0030	0031					
MOD	0044	0048										
NLA	0010	0014										
NLS	0012	0014	0046	0050								
NOL	0012	0014	0020	0021	0040	0046	0050					
NOS	0019	0020	0022	0041								
NSA	0011	0014										
NSE	0012	0014	0019	0046	0050							
NSS	0012	0014	0019	0046	0050							
SSQ	0016	0031	0031	0033								
SUM	0015	0030	0030	0032								
CINT	0037	0042	0046									
FPIC	0014											
IOPT	0004	0012	0034									
NTOT	0020	0032	0033	0033	0046	0050						
RAVE	0032	0033	0046	0050								
RFAC	0012	0014	0036	0037								
RMAX	0018	0027	0029	0036	0046	0050						
RMIN	0017	0024	0026	0036	0042	0046	0050					
RSTD	0033	0046	0050									
SQRT	0033											
FLOAT	0032	0033	0033	0037								
ICASE	0009	0013	0013	0044	0048							
IHIST	0002	0039	0043	0043	0046							
TARG ID	0003	0012	0046	0050								

LABEL	DEFINED	REFERENCES
1	0005	0012
9	0006	0044 0048
10	0007	0046
11	0008	0050
50	0012	0047 0051
55	0027	0024
59	0030	0027
60	0031	0021 0022
65	0039	0038
70	0043	0040 0041
80	0048	0034
100	0052	0012

LABEL	ADDR
50	000DF6
65	000FFA

LABEL	ADDR
55	000ED8
70	001054

LABEL	ADDR
59	000EE4
80	001144

LABEL	ADDR
60	000EEE
100	001104

\*STATISTICS\*      SOURCE STATEMENTS =      52 , PROGRAM SIZE =      4608

Figure A-1 (Cont'd)

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1 * THIS SUBROUTINE WAS CREATED TO PROCESS THE MARS MARINER TV DATA . IT 00000100
2 * BASICALLY COPIES A TAPE FILE TO A DISK ON THE FIRST CALL NAD 00000200
3 * THEN ACCESSES THE DISK IN DIRECT MODE TO FIND SQUARES OF THE 00000300
4 * PICTURE AS REQUESTED BY THE CALLING PROGRAM 00000400
5 PRINT NOGEN 00000500
6 FPIC CSECT 00000600
7 USING *,15 00000700
8 SAVE (14,12),,FPIC_DON_SPARROW 00000800
14 LR 14,13 CHAIN THE SAVE AREAS 00000900
15 LA 13,SAVE 00001000
16 ST 13,8(14) 00001100
17 ST 14,4(13) 00001200
18 B FPIC1 00001300

000000 000003 20 DROP 15 00001500
00001A 18ED 21 USING SAVE,13 00001600
00001C 4100 F030 22 SAVE DS 9D 00001700
000020 50DE 0008 23 TDCB DCB DDNAME=TAPE,DSORG=PS,MACRF=(R),EODAD=EOT,SYNAD=TERR, *00001800
000024 50ED 0304 74 DDCB DCB DDNAME=DISK,DSORG=PS,MACRF=(WL),EODAD=EOD,SYNAD=DERR, *00002000
000028 47F0 F58C 125 DDCBI DCB DDNAME=DISK,DSORG=DA,MACRF=(RKC),EODAD=EODI,SYNAD=DERRI,*00002200
000180 174 KEY DS F 00002300
000184 175 BUFFER DS CL1000 00002400
00056C 00 176 TRUE EQU 1 00002500
000570 177 FALSE EQU 0 00002600
000578 178 ENDOFHDS DC ALL(FALSE) 00002700
00057C 179 TEMP DS D 00002800
000580 180 NUMBEROFLINESPERPICTURE DS F 00002900
000584 181 NUMBEROFSAMPLESPERLINE DS F 00003000
000588 42000000 182 NOOFLABELRECORDS DS F 00003100
183 NUMBEROFRECORDSONTAPE DS F 00003200
184 FLOATPAT DC X'42000000' 00003300
00058C 186 FPIC1 DS OH 00003400
00058E 187 LR 8,1 SAVE THE PARM ADDRESS 00003500
000592 4780 D53C 188 CLI ENDOFHDS,TRUE FIRST ENTRY ? 00003600
000594 4780 D6FE 189 BE FPIC2 NO - BRANCH 00003700
0005A2 4110 D048 191 OPEN (TDCB,(INPUT)) YES - OPEN THE INPUT TAPE 00003800
0005A6 9110 1030 197 USING IHADC8,1 00003900
0005AA 47E0 D836 198 LA 1,TDCB 00004000
0005AE 4140 D0A0 199 TM DCOFLGS,X'10' WAS OPEN ALRIGHT ? 00004100
0005B2 1822 200 BNO TDERR NO - BRANCH 00004200
0005B4 47F0 D596 202 LA 4,DDCB YES - GET THE @ OF THE DISK DCB 00004300
0005B8 203 DROP 1 00004400
0005C6 205 SR 2,2 CLEAR THE RECORD COUNTER REGISTER 00004500
0005C8 206 B READTAPE START THE COPY 00004600
0005D8 208 COPYTODISK DS OH 00004700
0005DA 209 CHECK CDDECB SEE IF DISK WRITE WAS ALRIGHT 00004800
0005E6 215 READTAPE DS OH 00004900
0005E8 D703 D5A0 D5A0 005D0 005D0 216 XC TDECB,TDECB CLEAR THE DECB FOR A READ 00005000
0005EE 4122 0001 217 READ TDECB,SF,TDCB,BUFFER+4,'S' READ A TAPE RECORD 00005100
0005F2 9501 D53C 230 LA 2,1(2) COUNT THE RECORD 00005200
0005F4 4780 D624 232 CLI ENDOFHDS,TRUE ARE WE IN THE DATA RECORDS ? 00005300
0005F6 233 BE DATARECORD YES - BRANCH 00005400
000608 4110 D158 235 CHECK TDECB NO - SEE IF THE TAPE READ WAS ALRIGHT 00005500
00060C 5920 D938 240 LA 1,BUFFER+4 SET BASE FOR USING 00005600
000610 4720 D600 241 C 2,='1' IS THIS THE FIRST LABEL REOCD ? 00005700
000614 F273 D540 1020 00570 0000 242 BH NOTISTLABEL NO - BRANCH 00005800
00061A 4F00 D540 244 USING LABEL,1 YES - GET NS AND NL FOR THIS TAPE 00005900
00061E 5000 D548 245 PACK TEMP,NL CONVERT FROM EBCDIC TO BINARY 00006000
000622 F273 D540 1024 00570 00024 246 CVB 0,TEMP 00006100
000628 4F00 D540 247 ST 0,NUMBEROFLINESPERPICTURE 00006200
00062C 5000 D54C 0057C 249 PACK TEMP,NS REPEAT FOR NS 00006300
250 CVB 0,TEMP 00006400
251 ST 0,NUMBEROFSAMPLESPERLINE 00006500
252 DROP 1 00006600
253 00006700
254 * THE LAST LOGICAL LABEL RECORD HAS AN 'L' IN THE LAST BYTE - 00006800
255 * NO ADDITIONAL LABEL RECORDS WILL FOLLOW THIS ONE 00006900
000630 256 NOTISTLABEL DS OH 00007000
000630 4100 0005 257 LA 0,5 00007100
000632 00007200

```

Figure A-2 SUBROUTINE FPIC LISTING

000634			259	CHECKFORLASTLABEL DS OH		00008400
		00000	260	USING LABELREC,1		00008500
000634 9503 1047	00047		261	CLI ENDOFLABELCHECK,C*L' IS THIS THE LAST LABEL RECORD ?		00008600
000638 4780 D618		00648	262	BE SETLASTLABELRECORD		00008700
00063C 4111 0048		00048	264	LA 1,72(1) NO - BUMP INDEX		00008900
000640 4600 D604		00634	265	BCT 0,CHECKFORLASTLABEL LOOK ALL LABEL RECORDS IN BLOCK		00009000
000644 47F0 D596		005C6	266	B READTAPE GET NEXT LABEL RECORD		00009100
000648			268	SETLASTLABELRECORD DS OH		00009300
000648 9201 D53C	0056C		269	MVI ENDOFHDRS,TRUE YES - SET SWITCH		00009400
00064C 5020 D550		00580	270	ST 2,NDOFLABELRECORDS SAVE THE NUMBER OF LABEL RECORDS		00009500
000650 47F0 D596		005C6	271	B READTAPE CONTINUE		00009600
			272	DROP 1		00009700
000654			274	DATARECORD DS OH		00009900
			275	CHECK TDEC8 SEE IF TAPE READ WAS ALRIGHT		00010000
000662 9110 4030	00030		281	USING IHADCR,4		00010200
000666 4710 D656		00686	282	TM DCBOFLGS,X*10' IS THE DISK DCB OPEN ?		00010300
			283	BO DDCBO YES - BRANCH		00010400
00066A 5800 D54C		0057C	285	L 0,NUMBEROFSAMPLESPERLINE NO - SET THE BLKSIZE		00010600
00066E 4000 403E		0003E	286	STH 0,DCBBLKSI		00010700
00067E 9110 4030	00030		288	OPEN (DDCB,(OUTPUT)) OPEN THE DCB		00010900
000682 47E0 D846		00876	294	TM DCBOFLGS,X*10' WAS THE OPEN ALRIGHT ?		00011000
			295	BNO DOERR NO - BRANCH		00011100
000686			297	DDCBO DS OH		00011300
			298	* THE KEY OF EACH RECORD IS THE RECORD NUMBER OR LINE NUMBER WHICH		00011400
			299	* RANGES FROM 1 TO NUMBEROFLINESPERPICTURE		00011500
000686 1802			300	LR 0,2 COMPUTE THE KEY		00011600
000688 5800 D550		00580	301	S 0,NDOFLABELRECORDS		00011700
00069C 5000 D154		00184	302	ST 0,BUFFER SET THE KEY IN FRONT OF THE RECORD		00011800
000690 D703 D66C D66C 0069C 0069C			304	XC CODECB,CODECB CLEAR THE DCB FOR DISK WRITE		00012000
			305	WRITE CODECB,SF,DDCB,BUFFER WRITE THE RECORD TO DISK		00012100
0006BA 47FF D68E		0068E	319	B **4(15) CHECK THE RETURN CODE FROM WRITE		00012300
0006BE 47F0 D588		00588	320	B COPYTODISK		00012400
0006C2 47F0 D588		00588	321	B COPYTODISK		00012500
0006C6 47F0 D588		00588	322	B COPYTODISK		00012600
			324	ABEND 69,DUMP WRITE WAS NO GOOD		00012800
			332	DROP 4		00012900
0006DA			334	EOT DS OH		00013100
0006DA 0620			335	BCTR 2,0 (2) = NUMBER OF RECORDS ON TAPE		00013200
0006DC 5020 D554		00584	336	ST 2,NUMBEROFRECORDSONTAPE		00013300
0006E0 5820 D550		00580	338	S 2,NDOFLABELRECORDS DOES IT CHECK WITH NL IN HEADER ?		00013500
0006E4 5920 D548		00578	339	C 2,NUMBEROFLINESPERPICTURE		00013600
0006E8 4780 D6CA		006FA	340	BE TOOK YES - BRANCH		00013700
			342	ABEND 169,DUMP		00013900
0006FA			351	TOOK DS OH		00014100
			352	CLOSE (TDCB,,DDCB) CLOSE THE DCBS		00014200
00070A 4120 D0E8		00118	361	LA 2,DDCBI SET DATA TO READ DISK		00014400
		00000	362	USING IHADCB,2		00014500
00070E 5800 D54C		0057C	364	L 0,NUMBEROFSAMPLESPERLINE		00014700
000712 4000 2052		00052	365	STH 0,DCBLRECL		00014800
000716 4000 203E		0003E	366	STH 0,DCBBLKSI		00014900
			368	OPEN (DDCBI,(INPUT)) READY DISK FOR INPUT MODE		00015100
000726 9110 2030	00030		375	TM DCBOFLGS,X*10' WAS RE-OPEN ALRIGHT ?		00015300
00072A 47E0 D856		00886	376	BNO DOERR NO - BRANCH		00015400
			377	DROP 2		00015500
			379	* WE END UP AT THIS POINT AFTER THE TAPE IS COPIED TO THE DISK ON THE		00015700
			380	* FIRST CALL, OK IMMEDIATELY ON THE SECOND AND SUBSEQUENT CALLS		00015800
00072E			381	FPIC2 DS OH		00015900
		00000	382	USING PARMLIST,8		00016000
00072E 9825 8000		00000	383	LM 2,5,NLS GET THE ADDRESSES OF THE PARMS		00016100
000732 5822 0000		00000	384	L 2,0(2)		00016200
000736 5833 0000		00000	385	L 3,0(3)		00016300
00073A 5844 0000		00000	386	L 4,0(4)		00016400
00073E 5855 0000		00000	387	L 5,0(5)		00016500
000742 1222			389	LTR 2,2 IS START LINE NUMBER POSITIVE ?		00016700
000744 47D0 D986		00886	390	BNP A269 NO - BRANCH		00016800
000748 1812			392	LR 1,2		00017000
00074A 1A15			393	AR 1,5		00017100
00074C 5910 D548		00578	394	C 1,NUMBEROFLINESPERPICTURE IS END LINE # IN PICTURE ?		00017200
000750 4720 D896		008C6	395	BH A369 NO - BRANCH		00017300
000754 1233			397	LTR 3,3 IS THE START SAMPLE POSITIVE ?		00017500
000756 47D0 D8A6		008D6	398	BNP A469 NO - BRANCH		00017600
00075A 1934			400	CR 3,4 IS START SAMPLE LESS THAN END SAMPLE ?		00017800
00075C 47B0 D8B6		008E6	401	BNL A569 NO - BRANCH		00017900
000760 5940 D54C		0057C	403	C 4,NUMBEROFSAMPLESPERLINE IS END SAMPLE ON A LINE ?		00018100
000764 4720 D8C6		008F6	404	BH A669 NO - BRANCH		00018200

Figure A-2 (Cont'd)

000768 1244		406	LTR	4,4	IS END SAMPLE WITHIN A LINE ?	00018400
00076A 4700 D8D6	00906	407	BNP	A769	NO - BRANCH	00018500
00076E 1255		409	LTR	5,5	IS THE NUMBER OF LINES IN SQUARE POSITIVE	00018700
000770 4700 D8E6	00916	410	BNP	A869	NO - BRANCH	00018800
000774 5860 8014	00014	412	L	6,NLA	GET THE DIMENSIONS OF RECEIVING ARRAY	00019000
000778 5866 0000	00000	413	L	6,0(6)	# OF COLUMNS IN RECEIVING ARRAY	00019100
00077C 5870 8018	00018	414	L	7,NSA		00019200
000780 5877 0000	00000	415	L	7,0(7)	# OF ROWS, E.G. ARRAY 'A' IS 'A(NSA,NLA)'	00019300
000784 1814		417	LR	1,4		00019500
000786 1813		418	SR	1,3		00019600
000788 4111 0001	00001	419	LA	1,1(1)	(1) IS THE NUMBER OF SAMPLES TO PUT IN A	00019700
		420 *			PUT IN A ROW OF 'A' ARRAY	00019800
00078C 1917		421	CR	1,7	IS THE REQUEST FITTABLE IN ARRAY ROWS ?	00019900
00078E 4720 D8F6	00926	422	BH	A969	NO - BRANCH	00020000
000792 1956		424	CR	5,6	IS REQUEST FITTABLE IN ARRAY COLUMNS ?	00020200
000794 4720 D906	00936	425	BH	A1069	NO - BRANCH	00020300
000798 5810 801C	0001C	427	L	1,RFAC	GET THE SCALE FACTOR	00020500
00079C 7821 0000	00000	428	LE	2,0(1)		00020600
0007A0 58C0 8010	00010	430	L	12,A		00020800
0007A4 18A7		431	LR	10,7		00020900
0007A6 89A0 0002	00002	432	SLL	10,2	NSA*4 - SIZE OF ROW IN ARRAY	00021000
0007AA 181A		434	LR	1,10		00021200
0007AC 0650		435	BCTR	5,0		00021300
0007AE 1C05		436	MR	0,5		00021400
0007B0 4181 C000	00000	437	LA	11,0(1,12)	A + ((NSA-1)*4)*(NOL-1) - # OF LAST ROW IN	00021500
		438 *			ARRAY TO SET FROM DATA	00021600
0007B4 0630		440	BCTR	3,0	READY (3) FOR LOOP AT SAMPLELOOP	00021800
0007B6 4160 0001	00001	442	LA	6,1	INCREMENT FOR SAMPLE LOOP	00022000
0007BA 0640		443	BCTR	4,0		00022100
0007BC 4174 D154	00184	444	LA	7,BUFFER(4)	# OF LAST SAMPLE IN RECORD TO PROCESS	00022200
0007C0		446 LINELOOP	DS	0H		00022400
0007C0 5020 D150	00180	447	ST	2,KEY	SET OF RECORD WE WANT TO READ	00022500
0007C4 D703 D7A0 D7A0 007D0 007D0		449	XC	FDECB,FDECB	READY FOR FETCH READ	00022700
		450	READ	FDECB,DK,DDCBI,BUFFER,'S',KEY,TEMP	FETCH DATA FROM DISK	00022800
		464	CHECK	FDECB	SEE IF THE READ WAS ALRIGHT ?	00022900
000800 1811		470	SR	1,1	(1) IS INDEX TO ROW IN A ARRAY	00023100
000802 4143 D154	00184	471	LA	4,BUFFER(3)	(4) IS PTR TO DATA ITEM TO PROCESS	00023200
000806		473 SAMPLELOOP	DS	0H		00023400
000806 4304 0000	00000	474	IC	0,0(4)	GET THE DATA	00023500
00080A 4200 D559	00589	475	STC	0,FLOATPAT+1	FLOAT IT	00023600
00080E 7800 D558	00588	476	LE	0,FLOATPAT		00023700
000812 3002		477	DER	0,2	SCALE IT	00023800
000814 7001 C000	00000	478	STE	0,0(1,12)	STORE IT IN 'A' ARRAY	00023900
000818 4111 0004	00004	479	LA	1,4(1)	BUMP ROW PTR	00024000
00081C 8746 D7D6	00806	480	BXLE	4,6,SAMPLELOOP	FINISH THE ROW	00024100
000820 4122 0001	00001	482	LA	2,1(2)	BUMP DESIRED RECORD KEY	00024300
000824 87CA D790	007C0	483	BXLE	12,10,LINELOOP	PROCESS ALL LINES IN SQUARE	00024400
000828 58D0 0004	00004	485	L	13,4(13)		00024600
		486	RETURN	(14,12),RC=0	RETURN	00024700
000836		491 EOD	DS	0H		00024900
		492	ABEND	1169,DUMP		00025000
000846		501 DERR	DS	0H		00025200
		502	ABEND	1269,DUMP		00025300
000856		511 TERR	DS	0H		00025500
		512	ABEND	1369,DUMP		00025600
000866		521 TOERR	DS	0H		00025800
		522	ABEND	1469,DUMP		00025900
LOC OBJECT CODE	ADDR1 ADDR2	STMT	SOURCE STATEMENT			ASM H V 05 11.17 07/14/77
000876		531 DOERR	DS	0H		00026100
		532	ABEND	1569,DUMP		00026200
000886		541 DOUERR	DS	0H		00026400
		542	ABEND	1669,DUMP		00026500
000896		551 DERRI1	DS	0H		00026700
		552	ABEND	1769,DUMP		00026800
0009A6		561 DERRI2	DS	0H		00027000
		562	ABEND	1869,DUMP		00027100

Figure A-2 (Cont'd)

0008B6	571 A269	DS OH	00027300
	572	ABEND 269,DUMP	00027400
0008C6	581 A369	DS OH	00027600
	582	ABEND 369,DUMP	00027700
0008D6	591 A469	DS OH	00027900
	592	ABEND 469,DUMP	00028000
0008E6	601 A569	DS OH	00028200
	602	ABEND 569,DUMP	00028300
0008F6	611 A669	DS OH	00028500
	612	ABEND 669,DUMP	00028600
000906	621 A769	DS OH	00028800
	622	ABEND 769,DUMP	00028900
000916	631 A869	DS OH	00029100
	632	ABEND 869,DUMP	00029200
000926	641 A969	DS OH	00029400
	642	ABEND 969,DUMP	00029500
000936	651 A1069	DS OH	00029700
	652	ABEND 1069,DUMP	00029800
000946	661 EODI	DS OH	00030000
	662	ABEND 1969,DUMP	00030100
000956	671 DERRI	DS OH	00030300
	672	ABEND 2069,DUMP	00030400
000000	681 LABEL	DSECT	00030600
000000	682	DS 32X	00030700
000020	683 NL	DS CL4	00030800
000024	684 NS	DS CL4	00030900
000028	685	DS 32X	00031000
000000	687 LABELREC	DSECT	00031200
000000	688	DS 71X	00031300
000047	689 ENDOFLABELCHECK	DS C	00031400
000000	691 PARMLIST	DSECT	00031600
000000	692 NLS	DS A	00031700
000004	693 NSS	DS A	00031800
000008	694 NSE	DS A	00031900
00000C	695 NOL	DS A	00032000
000010	696 A	DS A	00032100
000014	697 NLA	DS A	00032200
000018	698 NSA	DS A	00032300
00001C	699 RFAC	DS A	00032400
	701	DCBD DSORG=PS	00032600
	702**,**	IMB069 DEVD NOT SPECIFIED-ALL ASSUMED	02-IMBEX
	1452	END	00032700
	1453	=F*1*	

000968 00000001

#### CROSS REFERENCE

SYMBOL	LEN	VALUE	DEFN	REFERENCES
A	00004	000010	0696	0430
A1069	00002	000936	0651	0425
A269	00002	0008B6	0571	0390
A369	00002	0008C6	0581	0395
A469	00002	0008D6	0591	0398
A569	00002	0008E6	0601	0401
A669	00002	0008F6	0611	0404
A769	00002	000906	0621	0407
A869	00002	000916	0631	0410
A969	00002	000926	0641	0422
BUFFER	01000	000184	0175	0225 0240 0302 0313 0444 0458 0471
CODECB	00004	00069C	0308	0210 0304 0304
CHECKFORLASTLABEL	00002	000634	0259	0265
COPYTODISK	00002	000588	0208	0320 0321 0322
DATARECORD	00002	000654	0274	0233
DCBBIT0	00001	00000080	0714	0835 0851 0890 0908 0953 0963 0979 1020 1030 1042 1065 1096 1099 1101 1103
DCBBIT1	00001	00000040	0715	1126 1129 1149 1154 1173 1210 1263 1286 1318 1322 1335 1430 1433 1443
DCBBIT2	00001	00000020	0716	0936 0852 0909 0917 0953 0963 0981 1021 1031 1044 1066 1068 1078 1096 1099
DCBBIT3	00001	00000010	0717	1102 1103 1131 1149 1152 1154 1176 1177 1178 1213 1214 1263 1268 1324 1326
DCBBIT4	00001	00000008	0718	1338 1382 1430 1435 1444
				0837 0853 0910 0919 0963 0982 1022 1032 1045 1047 1048 1066 1068 1073 1079
				1096 1098 1105 1133 1156 1158 1181 1182 1183 1217 1218 1265 1291 1327 1343
				1385 1389 1430 1445
				0901 0838 0854 0924 0966 0983 1023 1045 1048 1050 1066 1081 1106 1136 1156
				1160 1185 1186 1187 1221 1222 1265 1293 1295 1297 1329 1344 1385 1390 1430
				0902 0855 0925 0969 0984 1024 1033 1082 1107 1137 1162 1168 1169 1190 1191
				1225 1226 1228 1229 1267 1300 1345 1385 1391

Figure A-2 (Cont'd)

DCBBIT5	00001	00000004	0719	0804	0856	0927	0970	0984	1025	1034	1083	1110	1112	1139	1162	1165	1166	1169
DCBBIT6	00001	00000002	0720	0806	0839	0857	0929	0971	0988	1026	1084	1086	1089	1110	1113	1140	1201	1202
DCBBIT7	00001	00000001	0721	0808	0840	0933	0973	0989	1027	1084	1087	1089	1115	1144	1206	1207	1244	1245
DCBBLKSI	00002	00003E	1352	0286	0366													
DCBFDAD	00008	000005	0735	0733														
DCBLRECL	00002	000052	1414	0365														
DCBOFLGS	00001	000030	1125	0199	0282	0294	0375											
DCBSSIO	00008	000000	0942	0945														
DCBWTID	00004	000000	0862	0877	0886													
DDCB	00004	000000	0078	0202	0292	0312	0358											
DDCBI	00004	000118	0130	0361	0372	0457	0462											
DDCBO	00002	000686	0297	0283														
DERR	00002	000846	0501	0111														
DERRI	00002	000956	0671	0161														
DOERR	00002	000876	0531	0295														
DOERRR	00002	000986	0541	0376														
ENDOFHORS																		
ENDOFHORS	00001	00056C	0178	0188	0232	0269												
ENDOFLABELCHECK																		
ENDOFLABELCHECK	00001	000047	0689	0261														
EOO	00002	000836	0491	0096														
EOOI	00002	000946	0661	0146														
EOT	00002	0006DA	0334	0045														
FALSE	00001	00000000	0177	0178														
FDECB	00004	0007D0	0453	0449	0449	0465												
FLOATPAT	00004	000588	0184	0475	0476													
FPIC1	00002	00058C	0186	0018														
FPIC2	00002	00072E	0381	0189														
IMADCB	00001	00000000	0712	0197	0281	0362	0760	0784	0812	0831	0861	0941	1002	1053	1122	1254	1271	1275
KEY	00004	000180	0174	0447														
LABEL	00001	00000000	0681	0244														
LABELREC	00001	00000000	0687	0260														
LINELOOP	00002	0007C0	0446	0483														
NL	00004	000020	0683	0245														
NLA	00004	000014	0697	0412														
NLS	00004	000000	0692	0383														
NOOFLABELRECORDS																		
NOOFLABELRECORDS	00004	000580	0182	0270	0301	0338												
NOT1STLABEL																		
NOT1STLABEL	00002	000630	0256	0242														
NS	00004	000024	0684	0249														
NSA	00004	000018	0698	0414														
NUMBEROFLINESPERPICTURE																		
NUMBEROFLINESPERPICTURE	00004	000578	0180	0247	0339	0394												
NUMBEROFRECORDSONTAPE																		
NUMBEROFRECORDSONTAPE	00004	000584	0183	0336														
NUMBEROFSAMPLESPERLINE																		
NUMBEROFSAMPLESPERLINE	00004	00057C	0181	0251	0285	0364	0403											
PARMLIST	00001	00000000	0691	0382														
READTAPE	00002	0005C6	0215	0206	0266	0271												
RFAC	00004	00001C	0699	0427														
SAMPLELOOP																		
SAMPLELOOP	00002	000806	0473	0480														
SAVE	00008	000030	0022	0015	0021													
SETLASTLABELRECORD																		
SETLASTLABELRECORD	00002	000648	0268	0262														
TOCB	00004	000078	0027	0195	0198	0224	0356											
TOCB	00004	0005D0	0220	0216	0216	0236	0276											
TOCK	00002	0006FA	0351	0340														
TEMP	00008	000570	0179	0245	0246	0249	0250	0461										
TERR	00002	000856	0511	0060														
TOERR	00002	000866	0521	0200														
TRUE	00001	00000001	0176	0188	0232	0269												
=F*1*	00004	000968	1453	0241														

# DIAGNOSTIC CROSS REFERENCE AND ASSEMBLER SUMMARY

NO STATEMENTS FLAGGED IN THIS ASSEMBLY

OVERRIDING PARAMETERS- LOAD

OPTIONS FOR THIS ASSEMBLY

MODECK, OBJECT, LIST, XREF(SHORT), NORENT, NOTEST, BATCH, ALIGN, ESD, WORLD, LINECOUNT(55), FLAG(0), SYSPARM()

NO OVERRIDING DD NAMES

327 CARDS FROM SYSIN  
392 LINES OUTPUT

5496 CARDS FROM SYSLIB  
30 CARDS OUTPUT

Figure A-2 (Cont'd)

NSE: is the last sample within the line to be included in the  
retrieved data

NOL: is the number of consecutive lines for retrieval of data  
(including NLS)

These parameters define the rectangular area within the image.

ARRAY: is the array into which the retrieval data is put

NLA: is the second dimension of the array (the number of  
rows (lines) the array contains)

NSA: is the first dimension of the array (the number of columns  
(samples) the array contains)

These parameters define the array to receive the data.

RFAC: is a scale factor by which each sample is divided

Note: NLS, NSS, NSE, NOL, NLA, NSA are assumed to be INTEGER \*4  
arguments while ARRAY and RFAC are assumed to REAL \*4  
arguments and further, a previous statement:

DIMENSION ARRAY (NSA, NLA) is assumed to have been included in  
the program that uses FPIC, which is FORTRAN compatible.

The data to be manipulated by FPIC is assumed to reside on a tape  
defined by a DD card as follows:

```
//GO.TAPE DD DSN=name,DISP=OLD,UNIT=2400,  
// VOL=SER=volser,DCB=(RECFM=U,BLKSIZE=1000)  
// LABEL=(file #,SL,,IN)
```

As part of FPIC's operation, another DD card is required as follows:

```
//GO.DISK DD DSN=&&TEMP,DISP=NEW,UNIT=SYSDA,  
// DCB =(DSORG=DA,KEYLEN=4,SPACE =(TRK,(300,10))
```

The operation of FPIC is as follows:

- 1) The first call to the subroutine causes the data (not the label) records to be copied from the tape to the disk data set as a direct access file with the line number as the record key.
- 2) The required lines (records) are then read from the disk, the sample bytes are extracted, floated, scaled and stored in ARRAY.

Certain operational errors will cause user abends. These are listed below.

Abend	Reason
69	disk write rejected
169	no. of data records $\neq$ NL in header
269	start line number negative
369	requested number of lines $>$ NL in header
469	start sample # negative
569	start sample # $>$ end sample #
669	end sample # $>$ NS in header
769	end sample # negative
869	number of lines in square (NOL) negative
969	(end sample # - start sample #) $>$ NS in A array
1069	number of lines (NOL) $>$ NL in A array
1169	EODAD on disk (write)
1269	disk I/O error (write)
1369	Tape I/O error
1469	unable to open tape data set dcb



1569	unable to open disk data set (write) dcb (write)
1669	unable to open disk data set dcb (read)
1969	end of dataon disk data set (read)
2069	I/O error on disk data set (read)

A program was written to compute the average spectral reflectance within a particular spectral band given the spectral reflectivity of the sample and the spectral transmission of the filter under consideration. Figure A-3 is a listing of this program which was written in Fortran language. Each of the steps in the program is clearly identified by comments and no further explanation is required. As utilized in the present study, FIL (1,J) through FIL (6,J) represent the six spectral transmissions for the blue, green, red filters of Mariner VI and Marinter VII respectively. RBAR represent the average spectral reflectivity of the sample in each of these spectral bands and the value of RATIO (1) and RATIO (2) are the red-to-blue and red-to-green reflectivity ratios for the Mariner VI spectral bands while RATIO (3) and RATIO (4) are the corresponding values for the Mariner VII spectral bands.

```

C
C *****
C **
C ** THIS PROGRAM COMPUTES THE AVERAGE SPECTRAL REFLECTANCE OF A SAMPLE **
C ** FOR SPECTRAL BANDS GIVEN THE SPECTRAL TRANSMISSION AND THE **
C ** SAMPLE REFLECTANCE. **
C **
C *****
C
ISN 0002      DIMENSION FIL(6,151),R(151),SID(20),RBAR(6),RATIO(4)
ISN 0003      1 FORMAT(4(10X,F10.0))
ISN 0004      2 FORMAT(20A4)
ISN 0005      3 FORMAT(1H1,9X,20A4//15X 'AVERAGE REFLECTANCE'/ 20X 'BAND',2X,'1',
*8X,'2',8X,'3',8X,'4',8X,'5',8X,'6'// 20X, 6F9.3 //15X 'REFLECTANCE
* RATIO$', 10X, 4F8.3, 4(/))
ISN 0006      4 FORMAT( 10X,20A4//15X 'AVERAGE REFLECTANCE'/ 20X 'BAND',2X,'1',
*8X,'2',8X,'3',8X,'4',8X,'5',8X,'6'// 20X, 6F9.3 //15X 'REFLECTANCE
* RATIO$', 10X, 4F8.3, 4(/))
ISN 0007      CALL CLEAR(FIL(1,1),FIL(6,151))
C
C *****
C *
C * READ SPECTRAL TRANSMISSION DATA
C *
C *****
C
ISN 0008      READ(5,1) (FIL(I,J),J=2,151)
ISN 0009      READ(5,1) (FIL(2,J),J=2,151)
ISN 0010      READ(5,1) (FIL(3,J),J=2,151)
ISN 0011      READ(5,1) (FIL(4,J),J=1,151)
ISN 0012      READ(5,1) (FIL(5,J),J=1,151)
ISN 0013      READ(5,1) (FIL(6,J),J=1,151)
C
C *****
C *
C * NORMALIZE SPECTRAL TRANSMISSION DATA
C *
C *****
C
ISN 0014      DO 30 I=1,6
ISN 0015      S=0.0
ISN 0016      DO 22 J=1,151
ISN 0017      22 S=S+FIL(I,J)
ISN 0018      DO 24 J=1,151
ISN 0019      24 FIL(I,J)=FIL(I,J)/S
ISN 0020      30 CONTINUE
C
C *****
C *
C * READ SAMPLE SPECTRAL REFLECTANCE DATA
C *
C *****
C
ISN 0021      ICASE=0
ISN 0022      40 READ(5,2,END=70) SID
ISN 0023      ,ICASE=ICASE+1

```

Figure A-3 AVERAGE REFLECTIVITY MAIN PROGRAM LISTING

```

ISN 0024      READ(5,1) R
C
C *****
C *
C *   DETERMINE AVERAGE SAMPLE REFLECTANCE FOR EACH SPECTRAL BAND
C *
C *****
C
ISN 0025      DO 50 I=1,6
ISN 0026      RBAR(I)=0.0
ISN 0027      DO 45 J=1,151
ISN 0028      45 RBAR(I)=RBAR(I)+R(J)*FIL(I,J)
ISN 0029      50 CONTINUE
C
C *****
C *
C *   COMPUTE RATIOS OF AVERAGE REFLECTANCES
C *
C *****
C
ISN 0030      DO 55 N=1,4
ISN 0031      55 RATIO(N)=0.0
ISN 0032      IF(RBAR(1).GT.0.0) RATIO(1)=RBAR(3)/RBAR(1)
ISN 0034      IF(RBAR(2).GT.0.0) RATIO(2)=RBAR(3)/RBAR(2)
ISN 0036      IF(RBAR(4).GT.0.0) RATIO(3)=RBAR(6)/RBAR(4)
ISN 0038      IF(RBAR(5).GT.0.0) RATIO(4)=RBAR(6)/RBAR(5)
C
C *****
C *
C *   PRINT OUTPUT DATA AND RECYCLE TO PROCESS NEXT CASE, IF ANY.
C *
C *****
C
ISN 0040      IF(MOD(ICASE,5).NE.1) GO TO 60
ISN 0042      WRITE(6,3) (SID(K),K=1,20),(RBAR(I),I=1,6),(RATIO(N),N=1,4)
ISN 0043      GO TO 40
ISN 0044      60 WRITE(6,4) (SID(K),K=1,20),(RBAR(I),I=1,6),(RATIO(N),N=1,4)
ISN 0045      GO TO 40
C
C *****
C *
C *   TERMINATE EXECUTION
C *
C *****
C
ISN 0046      70 STOP
ISN 0047      END

```

Figure A-3 (Cont'd)

LABEL	ADDR	LABEL	ADDR	LABEL	ADDR	LABEL	ADDR
22	001472	24	00149C	30	0014AA	40	0014C6
45	00153A	50	001550	55	00155E	60	001626
70	00167A						

\*\*\*\*\*F O R T R A N   C R O S S   R E F E R E N C E   L I S T I N G\*\*\*\*\*

SYMBOL	INTERNAL STATEMENT NUMBERS													
I	0014	0017	0019	0019	0025	0026	0028	0028	0028	0042	0042	0042	0044	0044
	0044													
J	0008	0008	0008	0009	0009	0009	0010	0010	0010	0011	0011	0011	0012	0012
	0012	0013	0013	0013	0016	0017	0018	0019	0019	0027	0028	0028		
K	0042	0042	0042	0044	0044	0044								
N	0030	0031	0042	0042	0042	0044	0044	0044						
R	0002	0024	0028											
S	0015	0017	0017	0019										
FIL	0002	0007	0007	0008	0009	0010	0011	0012	0013	0017	0019	0019	0028	
MOD	0040													
SID	0002	0022	0042	0044										
RBAR	0002	0026	0028	0028	0032	0032	0032	0034	0034	0034	0036	0036	0036	
	0038	0038	0038	0042	0044									
CLEAR	0007													
ICASE	0021	0023	0023	0040										
RATIO	0002	0031	0032	0034	0036	0038	0042	0044						

LABEL	DEFINED	REFERENCES						
1	0003	0008	0009	0010	0011	0012	0013	0024
2	0004	0022						
3	0005	0042						
4	0006	0044						
22	0017	0016						
24	0019	0018						
30	0020	0014						
40	0022	0043	0045					
45	0028	0027						
50	0029	0025						
55	0031	0030						
60	0044	0040						
70	0046	0022						

Figure A-3 (Cont'd)